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14 **Chapter 3**  
15 **Investigation of solar irradiance variations and**  
16 **their impact on middle atmospheric ozone**

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19 **Abstract** The satellite spectrometer SCIAMACHY aboard ENVISAT is a unique in-  
20 strument that covers at a moderately high spectral resolution the entire optical range  
21 from the near UV (230 nm) to the near IR ( $2.4\mu\text{m}$ ) with some gaps above  $1.7\mu\text{m}$ .  
22 This broad spectral range allows not only the retrieval of several atmospheric trace  
23 gases (among them ozone), cloud and aerosol parameters, but also regular daily  
24 measurements of the spectral solar irradiance (SSI) with an unprecedented spectral  
25 coverage. The following studies were carried out with irradiance and ozone data  
26 from SCIAMACHY: a) SCIAMACHY SSI was compared to other solar data from  
27 space and ground as well as with SIM/SORCE (Solar Irradiance Monitor, the only  
28 other satellite instrument daily measuring the visible and near IR, in order to verify  
29 the quality of the SCIAMACHY measurements, b) an empirical solar proxy model,  
30 in short the SCIA proxy model, was developed that permits expressing the SCIA-

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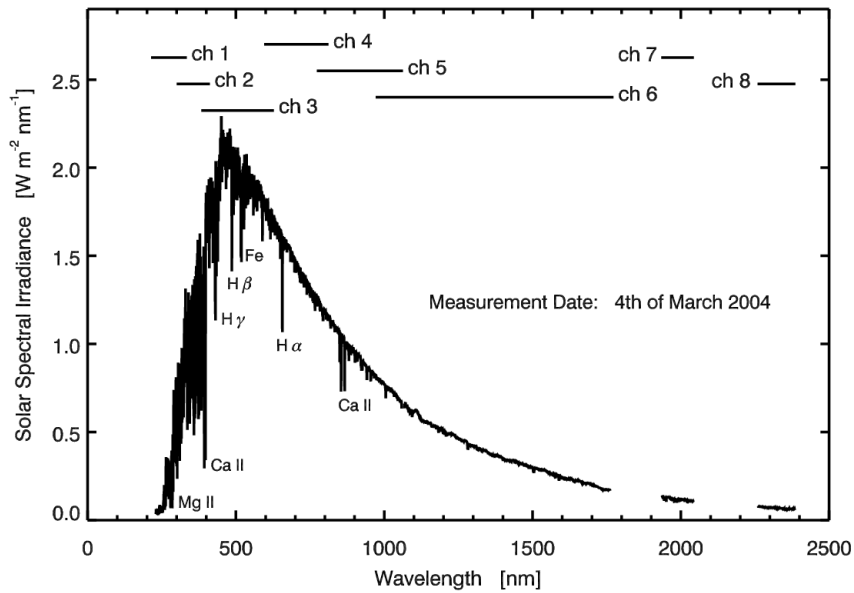
31 MACHY SSI variations by fitting solar proxies for faculae brightening and sunspot  
32 darkening, which then allows investigation of solar variability on time scales be-  
33 yond the instrument life time, e.g. 11-year solar cycle, c) solar cycle SSI variations  
34 derived from empirical models (Lean2000, SATIRE, SCIA proxy) and different ob-  
35 servations (SBUV composite, SUSIM) were compared for the three most recent  
36 solar cycles 21-23, and d) SCIAMACHY ozone limb profiles were analysed to de-  
37 rive signatures of the 27-day solar rotation on stratospheric ozone. Our studies were  
38 complemented by investigations of daytime variations in mesospheric ozone (here  
39 data from SABER/TIMED), which were compared to results from the HAMMO-  
40 NIA chemistry climate model.

### 41 3.1 Introduction

42 Regular daily space-borne satellite SSI monitoring started in 1978. The wavelength  
43 coverage of early SSI measurements from different satellite instruments was gen-  
44 erally limited to below 400 nm (UV), where the largest variations occur over an  
45 11-year solar cycle (Rottman *et al.*, 2004). A limiting factor for many space spec-  
46 trometers measuring in the UV is the optical degradation due to hard radiation that  
47 makes it challenging to maintain the accuracy over the instrument lifetime which  
48 rarely extends to more than a decade (DeLand *et al.*, 2004). In order to derive esti-  
49 mates for SSI variations over an entire 11-year solar cycle or more one needs to rely  
50 on a SSI timeseries composed of different instruments (UV composite) as done for  
51 the UV spectral range (DeLand and Cebula, 2008) or use solar proxies, like the Mg  
52 II index, that are well correlated with irradiance changes over a large spectral range  
53 to extrapolate beyond the instrumental lifetime (DeLand and Cebula, 1993; Viereck  
54 *et al.*, 2001).

55 Daily observations of the visible and near-IR started with the three channel SPM  
56 (Sun Photometer) of VIRGO/SOHO (1996-2010) at selected wavelength bands  
57 (Fröhlich *et al.*, 1997) and were continued with GOME/ERS-2 (Global Ozone  
58 Monitoring Experiment) since 1995, covering 240–800 nm (Weber *et al.*, 1998;  
59 Burrows *et al.*, 1999), SCIAMACHY/ENVISAT (Scanning Imaging Absorption  
60 Spectrometer for Atmospheric Chartography) since 2002, covering 220nm - 2.4 $\mu$ m  
61 (Bovensmann *et al.*, 1999), and SIM/SORCE (Solar Irradiance Monitor) since 2003,  
62 240nm–3 $\mu$ m (Harder *et al.*, 2005a,b). In Fig. 3.1 a sample SCIAMACHY solar ir-  
63 radiance spectrum is shown. Compared to the UV region, daily irradiance measure-  
64 ments simultaneously covering the UV, visible, and the near IR do not cover yet a  
65 complete solar cycle. One of the important scientific question is what are the irra-  
66 diance changes in the visible and near IR during 27-day solar rotations and can we  
67 use this information to extrapolate to changes during the 11-year solar cycle.

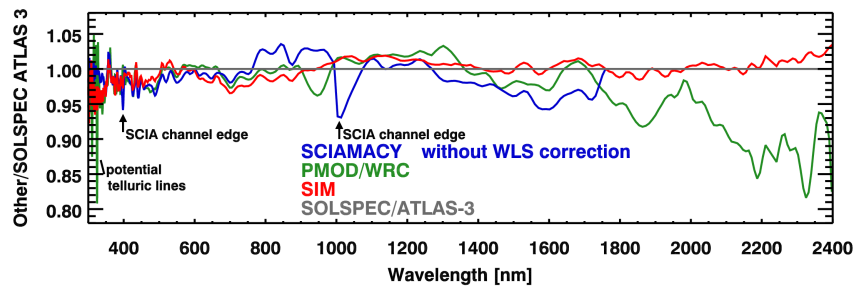
68 SCIAMACHY is primarily an atmospheric sounder measuring several trace  
69 gases in nadir (column amounts) and limb viewing geometry (vertical profiles)  
70 (Bovensmann *et al.*, 2011). Global vertical profiles are measured by SCIAMACHY  
71 and cover the altitude range from the tropopause to about 70 km altitude (*von Savi-*



**Fig. 3.1** SCIAMACHY full disc solar spectrum measured on March 4, 2004. The eight spectral channels varying in spectral resolution from 0.2 nm to 1.5 nm are indicated. The gaps near 1850 nm as well as 2200 nm are not measured by SCIAMACHY since terrestrial water vapor absorption saturates in the atmospheric observation mode. From *Pagaran et al. (2009)*. Reproduced by permission of the AAS.

72 *gny et al., 2005; Sonkaew et al., 2009*). The influence of irradiance variations related  
 73 to the 27-day mean solar rotation period on upper stratosphere ozone can be investi-  
 74 gated using SCIAMACHY ozone data. The upper stratosphere above 30 km is  
 75 chemically controlled and an immediate radiative influence on the photochemistry  
 76 is expected (e.g. *Gruzdev et al., 2009; Fioletov, 2009*). In this study for the first time  
 77 a wavelet analysis was applied to study the 27-day signature in ozone. This permits  
 78 the investigation of the time-varying frequency content of the ozone signal.

79 The non-polar orbit of the TIMED satellite (Thermosphere, Ionosphere, Meso-  
 80 sphere, Energetics and Dynamics) carrying the SABER instrument (Sounding of  
 81 the Atmosphere using Broadband Emission Radiometry) (*Russell III et al., 1999*)  
 82 permits the study of daytime variations in mesospheric ozone that are significantly  
 83 larger than the 27-day and solar cycle related changes observed in the upper strato-  
 84 sphere (*Huang et al., 2008*). In this study the daytime variation of mesospheric  
 85 ozone were compared for the first time with the output of a chemistry climate model  
 86 (*Dikty et al., 2010a*).



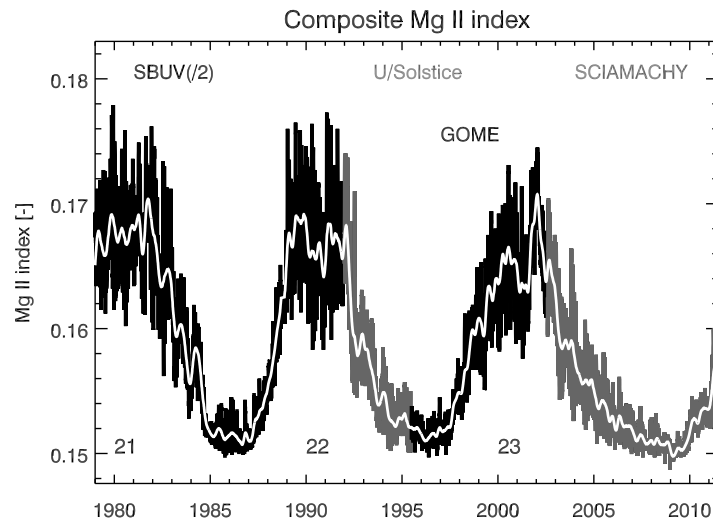
**Fig. 3.2** SSI ratios of SCIAMACHY, SIM (*Harder et al.*, 2010), and PMOD/WRC (WRC85) composite (*Wehrli*, 1985) with respect to the SOLSPEC/ATLAS-3 shuttle experiment data (*Thuillier et al.*, 2004). Close to the channel boundaries of the SCIAMACHY instrument larger deviations are observed due to instrumental artifacts. From *Pagaran et al.* (2011a). Reproduced with permission ©ESO.

### 87 3.2 SCIAMACHY spectral solar irradiance

88 SCIAMACHY is a passive remote sensing double spectrometer combining a predis-  
 89 persing prism and eight gratings in separate channels. Silicon and InGaAs detectors  
 90 are used as linear arrays with 1024 pixels each in Channels 1-5 (UV/visible) and  
 91 Channels 6-8 (near IR), respectively. A detailed description of SCIAMACHY can  
 92 be found in *Bovensmann et al.* (1999) and *Pagaran et al.* (2009).

93 Radiometrically calibrated SSI has been measured by SCIAMACHY since July  
 94 2002 once a day. A sample spectrum from March 2004 is shown in Fig. 3.1. The  
 95 SCIAMACHY SSI has been compared with solar data from other satellites and  
 96 measurements from the ground (*Skupin et al.*, 2005a,b; *Pagaran et al.*, 2011a).  
 97 Figure 3.2 shows the comparison of SCIAMACHY with SIM (*Harder et al.*,  
 98 2010), the SOLSPEC/ATLAS-3 shuttle experiment (*Thuillier et al.*, 2004), and the  
 99 PMOD/WRC (WRC85) composite (*Wehrli*, 1985). The PMOD/WRC composite  
 100 (200 nm – 10  $\mu$ m) was derived from various spectra obtained from aircraft, rocket,  
 101 and balloon experiments as well as ground data from *Neckel and Labs* (1984). SCIA-  
 102 MACHY agrees to within 5% (SIM within 4%) with the SOLSPEC data from 300 to  
 103 1600 nm (*Pagaran et al.*, 2011a). The theoretical precision is usually in the range of  
 104 2-3% based upon radiometric standards (*Bovensmann et al.*, 1999). A more compre-  
 105 hensive comparison also to other solar data can be found in *Pagaran et al.* (2011a).

106 In later years of the SCIAMACHY mission the optical degradation in the UV  
 107 due to the hard radiation environment in space is evident. The agreement of SCIA-  
 108 MACHY with other solar data can be improved when using the white light lamp  
 109 (WLS) source as a degradation correction, however, the corrections are too strong  
 110 since WLS itself is optically degrading and therefore this type of correction cannot  
 111 be applied to the more recent SCIAMACHY data (*Pagaran et al.*, 2011a). Further  
 112 investigations are underway to improve upon the in-flight radiometric calibration.  
 113 For atmospheric studies this is generally not a problem since the degradation cancels  
 114 out in the sun-normalized earth radiances used in most atmospheric retrievals.



**Fig. 3.3** Composite Mg II index measured near 280 nm derived from multiple satellite data and extending from solar cycle 21 to 24. In addition to SBUV(/2), UARS/Solstice, GOME, and SCIAMACHY as indicated here, other data such as SUSIM, the more recent SBUV/2 instruments from NOAA-16 to NOAA-18 as well as SORCE/Solstice have been used to fill daily gaps. The smooth white line shows the low-pass filtered time series by applying a 55-day triangular filter to remove the 27-day solar rotation signature.

115 The Mg II core to wing ratio derived from the Mg II Fraunhofer lines at 280  
 116 nm (Fig. 3.1) is an index that has been proven to correlate well with UV irradiance  
 117 changes down to 30 nm (*DeLand and Cebula, 1993; Viereck et al., 2001*). It  
 118 is a measure for the chromospheric activity of the sun and describes the plage and  
 119 faculae brightening responsible for the UV increase. The Mg II index, defined as  
 120 a ratio, is insensitive to instrumental degradation and has been derived from many  
 121 different instruments to provide a long-term time series going back to the late 1970s  
 122 (*DeLand and Cebula, 1993; Viereck and Puga, 1999; Viereck et al., 2004*). An up-  
 123 dated composite Mg II index by adding the GOME (*Weber et al., 1998; Weber,*  
 124 *1999*) and recent SCIAMACHY data is shown in Fig. 3.3. It seems that the Mg II  
 125 index was lower during the recent solar minimum in 2008 than the two solar minima  
 126 before, but this is not statistically significant. A potential lower solar minimum  
 127 value could be expected from the very low thermospheric density observed in 2008  
 128 (*Emmert et al., 2010*). Solar irradiance at extreme ultraviolet (EUV) wavelengths  
 129 heats the thermosphere, causing it to expand. Low EUV irradiance contracts the  
 130 thermosphere and decreases the density at a given altitude. The cooling of the upper  
 131 atmosphere due to increases of greenhouse gases can only explain part of the recent  
 132 contraction observed (*Emmert et al., 2010; Solomon et al., 2011*).

### 133 3.3 Irradiance variations from solar rotations to several solar 134 cycles

135 In order to estimate SSI irradiance variations beyond the instrument lifetime and  
136 covering several decades a model was developed that uses solar proxies scaled to  
137 SCIAMACHY SSI observations. The underlying assumption is that irradiance vari-  
138 ations are mainly caused by solar surface magnetic activity (*Fligge et al., 2000*) and  
139 can be expressed in terms of faculae brightening as represented by the Mg II index  
140 and sunspot darkening as expressed by the photometric sunspot index (PSI), here  
141 taken from *Balmaceda et al. (2009)*. The SSI can then be written as a time series as  
142 follows

$$I_{\lambda}(t) = I_{\lambda}(t_0) + a_{\lambda} [P_a(t) - P_a(t_0)] + b_{\lambda} [P_b(t) - P_b(t_0)] + p_{\lambda}(t), \quad (3.1)$$

143 where  $P_a(t)$  and  $P_b(t)$  are the Mg II index and PSI time series, respectively. A similar  
144 approach was used to model UV irradiance variations derived from UARS/Solstice  
145 (*Lean et al., 1997*).

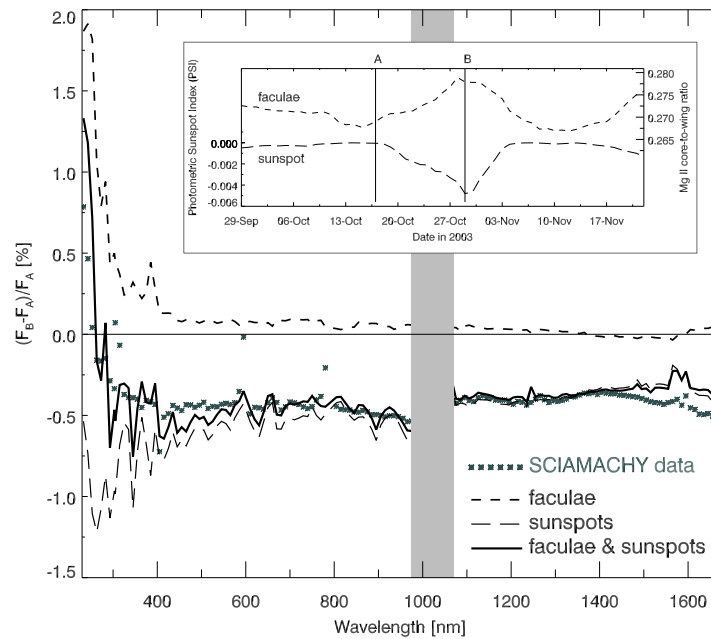
146 A multivariate linear regression is performed to determine the regression coef-  
147 ficients of the solar proxies. In addition to the two solar proxy terms piecewise  
148 polynomials,  $p_{\lambda}(t)$ , are used to correct for instrument degradation and small jumps  
149 following instrument and satellite platform anomalies (*Pagaran et al., 2009*). The  
150 regression was applied to SCIAMACHY SSI time series over several solar rotations  
151 during 2003 and 2004. Regression coefficients,  $a_{\lambda}$  and  $b_{\lambda}$ , were determined from  
152 240 nm to 1750 nm (SCIAMACHY channels 1 to 6) in steps of 10 nm (*Pagaran*  
153 *et al., 2009*). As a solar reference spectrum,  $I_{\lambda}(t_0)$ , the SCIAMACHY SSI from  
154 March 4, 2004, (Fig. 3.1) was selected.

155 The modeled and observed SCIAMACHY solar irradiance change is shown as  
156 an example in Fig. 3.4 during the Halloween 2003 solar storm, where the PSI index  
157 value reached the lowest value since 1980 and substantial mesospheric ozone loss  
158 (mainly due to solar protons) was observed by SCIAMACHY (*Rohen et al., 2005*).  
159 The combined faculae and sunspot contributions and SCIAMACHY observations  
160 are in qualitative agreement with Figure 6 in *Lean et al. (2005)*. Across the near-  
161 UV, vis, and near-IR spectral range solar irradiance dropped by 0.3% (near-IR) to  
162 0.5% (near-UV). This is consistent with a drop of about 0.4% in the total solar  
163 irradiance (TSI) or solar constant. Below 300 nm an irradiance enhancement due to  
164 faculae activity was observed reaching +1.3% near 250 nm.

165 The SCIAMACHY irradiance timeseries as well as the SCIA proxy model show  
166 the dark faculae effect in the spectral region 1400–1600 nm (near opacity  $H^{-}$  mini-  
167 mum), where both sunspot and faculae contributions are negative in agreement with  
168 observations from ground indicating a darkening under enhanced solar activity con-  
169 ditions (*Moran et al., 1992*). The SCIA proxy model, nevertheless, underestimates  
170 the observed irradiance depletion in this spectral region.

171 The SCIA proxy model can be used to reconstruct spectral irradiance changes  
172 since the late 1970s, where the Mg II index record started, covering nearly three  
173 solar cycles. From the SCIA proxy model the UV contribution below 400 nm to TSI



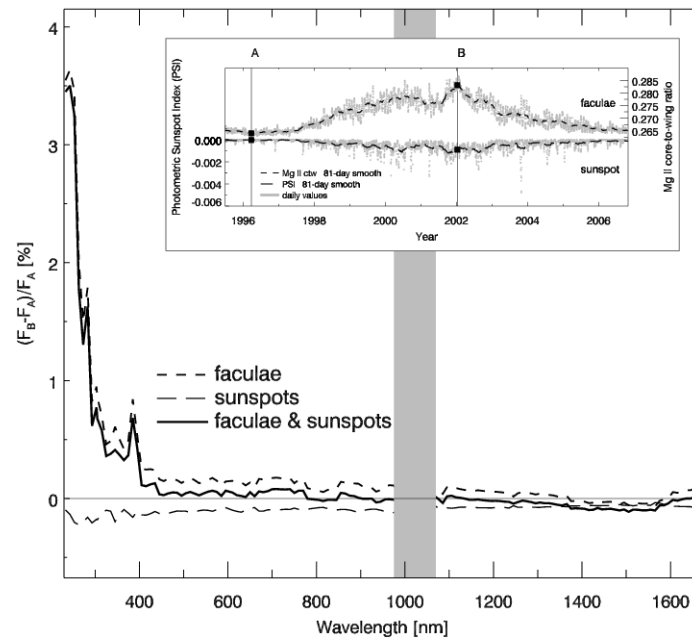


**Fig. 3.4** Modeled and observed SCIAMACHY solar irradiance change during the Halloween solar storm in 2003 decomposed into faculae and sunspot contributions. The inset shows the Mg II and PSI index with labels A and B indicating dates from which irradiance differences were derived. From Pagaran *et al.* (2009). Reproduced by permission of the AAS.

174 changes in solar cycle 23 ( $\sim 0.1\%$ ) is 55% (Pagaran *et al.*, 2009) which is higher  
 175 than the 30% estimate from solar cycle 22 derived from SOLSTICE observations  
 176 (Lean *et al.*, 1997) and lower than the 63% derived from the semi-empirical model  
 177 SATIRE (Spectral and Total Irradiance Reconstructions) (Krivova *et al.*, 2006).

178 The largest TSI change contribution comes from the near UV (300–400 nm),  
 179 where the irradiance solar cycle change per wavelength is well below 1% (Pagaran  
 180 *et al.*, 2009). During solar cycles 21 to 23, the dominant contribution to irradiance  
 181 changes in the UV from solar minimum to maximum comes from the faculae bright-  
 182 ening. The sunspot contribution is non-negligible in the near UV and in the visible  
 183 cancels within the error bars the faculae brightening (see Fig. 3.5, Pagaran *et al.*  
 184 (2009, 2011b)). The dark faculae near 1400–1600 nm are again evident at solar  
 185 maximum in agreement with observations by SIM and results from the SATIRE  
 186 model (Unruh *et al.*, 2008).

187 Harder *et al.* (2009) reported on SIM irradiance changes during the descending  
 188 phase of solar cycle 23 (April 2004 to November 2007) and found UV changes  
 189 that are much larger than models like the NRLSSI irradiance model (Lean, 2000)  
 190 indicate. This is also true when comparing to other data sets as shown in Fig. 3.6  
 191 where the comparison is extended to the SCIA proxy model, the SATIRE model  
 192 (Krivova *et al.*, 2009), and the UV composite from DeLand and Cebula (2008) as



**Fig. 3.5** Solar irradiance variations during solar cycle 23 as derived from SCIAMACHY observations and proxy data. Solar maximum and minimum dates were defined by the 81 day boxcar smooth of Mg II index timeseries (inset). Contributions from faculae and sunspots are indicated. From *Pagaran et al.* (2009). Reproduced by permission of the AAS.

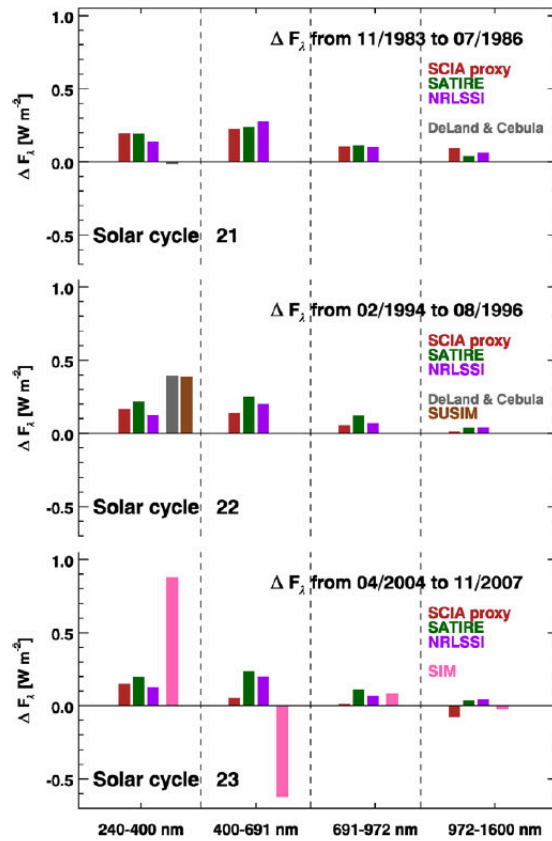
193 well. Also shown in this figure are the comparison of irradiance changes during the  
 194 descending phase of prior solar cycles with similar Mg II index change as in solar  
 195 cycle 23 (*Pagaran et al.*, 2011b).

196 It appears that current models including the SCIA proxy model that assume that  
 197 irradiance changes are mostly related to surface magnetic activity are underestim-  
 198 ating solar cycle changes in the UV as compared to the SIM observations. Direct  
 199 observations from SUSIM and the UV composite also see larger UV changes dur-  
 200 ing solar cycle 22 than the models, but are still only about half of SIM's result for  
 201 solar cycle 23. Such a large UV change as observed by SIM has strong implications  
 202 on radiative forcing in the upper atmosphere (*Haigh et al.*, 2010; *Oberländer et al.*,  
 203 2012) and will remain a matter of debate.

### 204 **3.4 Solar rotation (27-day) signature in stratospheric ozone**

205 The solar variation on the 11-year time scale has been shown to cause 2-3% variabil-  
 206 ity in tropical ozone at altitudes of approximately 40 km. This has been concluded  
 207 from different satellite observations (e.g. *Remsberg*, 2008; *Fioletov*, 2009, and ref-

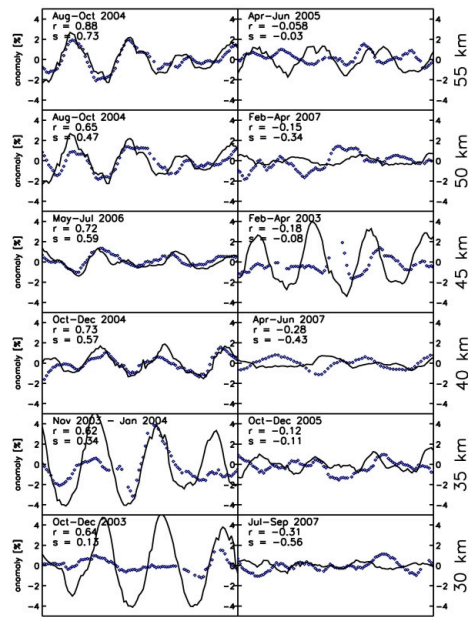
**Fig. 3.6** SSI changes during part of the descending phase of solar cycles 21 to 23 (top to bottom), respectively. Dates near solar maximum and minimum are chosen in such a way that the differences between the Mg II indices are about the same in each solar cycle and correspond to that of the SIM observation period used here. NRLSSI (*Lean, 2000*), SATIRE (*Krivova et al., 2009*) and SCIA proxy are models. SUSIM, SIM, and UV composite from *DeLand and Cebula (2008)* are direct satellite observations. From *Pagaran et al. (2011b)*.



208 erences therein) and was confirmed by model studies (e.g. *Langematz et al., 2005*;  
 209 *Sekiyama et al., 2006*; *Marsh et al., 2007*). The influence of the 27-day solar rota-  
 210 tion on ozone was first investigated by *Hood (1986)* in the 1980s using SBUV  
 211 ozone measurements. He found the ozone sensitivity at 45 km to be slightly more  
 212 than 0.4% per 1% change in the 205 nm flux. Further investigations with different  
 213 satellite data sets and model outputs covering other time periods followed (*Gruzdev*  
 214 *et al., 2009*; *Fioletov, 2009*, and references therein). *Austin et al. (2007)* and *Gruzdev*  
 215 *et al. (2009)* compared the 27-day ozone variability determined by chemistry cli-  
 216 mate models (CCM) with satellite measurements and were able to verify the obser-  
 217 vations in magnitude (0.4 to 0.5 %) but found the maximum ozone sensitivity  
 218 slightly lower in altitude (approx. 40 km) in the model simulations.

219 The motivation for this study is to use the new dataset that is available from  
 220 SCIAMACHY, e.g. global ozone profiles during the descending phase of solar cycle  
 221 23 (*von Savigny et al., 2005*; *Sonkaew et al., 2009*). Continuous wavelet transform  
 222 (CWT), fast Fourier transform (FFT), and cross correlations (CC) have been applied  
 223 to SCIAMACHY ozone in the tropics (<20° latitude) between 20 and 60 km altitude  
 224 (*Dikty et al., 2010b*). The maximum correlation between the Mg II index and ozone

**Fig. 3.7** Selected three month periods with high (left panels) and low correlation (right panels) between ozone (solid line) and Mg II index (circles). In each panel, the period, correlation ( $r$ ), and ozone sensitivity ( $s$ ) is indicated, the latter is defined as the ozone change per Mg II index change in units of  $\%/%$ . The ozone sensitivity per unit 205 nm solar irradiance change is obtained by multiplying  $s$  with 0.61. From *Dikty et al.* (2010b). Reproduced by permission of American Geophysical Union. ©2010 American Geophysical Union.



225 is weaker during the maximum of solar cycle 23 ( $r = 0.38$ ) than in the previous two  
 226 solar cycles that have been investigated in earlier studies using different data sets.  
 227 This is in agreement with results from *Fioletov* (2009).

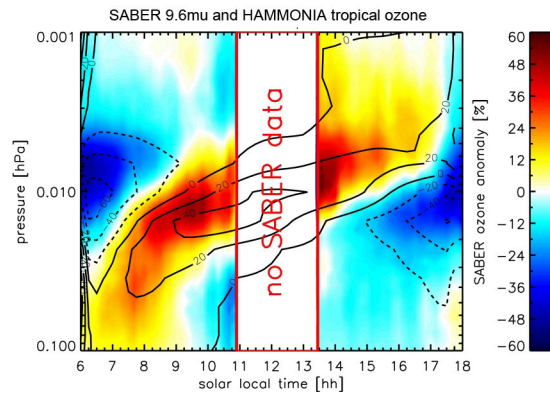
228 The magnitude of the ozone signals is highly time dependent as revealed by the  
 229 CWT analysis and may vanish for several solar rotations even close to solar maxi-  
 230 mum conditions (see Fig. 3.7). The ozone sensitivity (ozone change in percent per  
 231 percent change in 205 nm solar flux) is on average about  $0.2\%/%$  above 30 km  
 232 altitude and smaller by about a factor of two compared to earlier studies. For se-  
 233 lected three month periods the sensitivity may rise beyond  $0.6\%/%$  in better agree-  
 234 ment with earlier studies. The analysis of the 27-day solar forcing was also carried  
 235 out with stratospheric temperatures from the European Centre for Medium-Range  
 236 Weather Forecasts operational analysis. Although direct radiation effects on tem-  
 237 perature are weak in the upper stratosphere, temperature signals with statistically  
 238 significant periods in the 25–35 day range similar to ozone were found (*Dikty et al.*,  
 239 2010b).

### 240 3.5 Daytime variations in mesospheric ozone

241 In comparison with the 27-day solar rotation signal and the 11-year solar cycle re-  
 242 sponse in the stratosphere, the diurnal and daytime variation of UV radiation inflicts  
 243 a by far greater response in upper atmosphere ozone. The response of ozone above  
 244 60 km to variations in UV radiation is less well established. Ozone and temperature

245 data from SABER (Sounding of the Atmosphere using Broadband Emission Ra-  
 246 diometry) in its version 1.07 (*Russell III et al.*, 1999) are used to study the daytime  
 247 pattern of mesospheric ozone. In contrast to SCIAMACHY, SABER aboard TIMED  
 248 flies in a more inclined orbit allowing measurements at different local times. In our  
 249 study (*Dikty et al.*, 2010a) a specific sampling of SABER data was performed to de-  
 250 rive daytime pattern in tropical ozone using both the results from the  $1.27\mu$  air glow  
 251 (*Mlynczak et al.*, 2007) and  $9.6\mu$ m thermal emission retrieval (*Rong et al.*, 2008).  
 252 Compared to the earlier study on daytime variations by *Huang et al.* (2008) more  
 253 years of SABER data were used and our results were compared to HAMMONIA  
 254 (Hamburg Model of the Neutral and Ionized Atmosphere) (*Schmidt et al.*, 2006).

**Fig. 3.8** SABER observations at  $9.6\mu$ m (color) and HAMMONIA model (contour) daytime ozone variations between 0.1 and 0.001 hPa expressed as percent deviation from the daytime mean. From *Dikty et al.* (2010a).



255 The amplitude of daytime ozone variations is approximately 60% of the day-  
 256 time mean for SABER and lower for the model (see Fig. 3.8). The agreement  
 257 with HAMMONIA is generally better for the  $9.6\mu$ m retrieved ozone data than for  
 258 the  $1.27\mu$ m air glow retrieval (*Dikty et al.*, 2010a). The maximum daytime peak  
 259 anomaly observed at 0.05 hPa ( $\sim 70$  km) in the morning shifts its altitude to about  
 260 0.007 hPa ( $\sim 80$  km) in the afternoon. This daytime shift is in very good agreement  
 261 with the model, however the peak anomaly reaches a maximum of 40–50% of the  
 262 daytime mean, which is higher than HAMMONIA (30–40%). Negative anomalies  
 263 are observed in the early morning hours at 0.007 hPa and in the late afternoon near  
 264 0.015 hPa in quite good agreement with the model. During equinox the daytime  
 265 maximum ozone abundance is higher than during solstice, especially above 0.01 hPa  
 266 (approx. 80 km). The seasonal variation is somewhat weaker in HAMMONIA.

267 In contrast to ozone, temperature data from SABER (*Remsberg et al.*, 2008) show  
 268 little daytime variations between 65 and 90 km and their amplitudes are less than  
 269 1.5%, suggesting photochemistry playing a dominant role in the mesospheric ozone  
 270 chemistry. *Marsh et al.* (2002) proposed that the solar diurnal tide brings down  
 271 atomic oxygen for ozone production in the afternoon ( $> 85$  km). The model, how-  
 272 ever, underestimates ozone in the afternoon above approximately 0.01 hPa, so the  
 273 remaining difference could be attributed to solar tides. The minimum early in the  
 274 morning is caused by the direct photolysis of ozone before enough atomic oxy-

275 gen is produced to counteract the ozone destruction. The ozone rise in the morning  
276 hours may also be due to tides transporting ozone rich air from below (*Marsh et al.*,  
277 2002). The new SABER version 1.08 data will also include water vapor, which will  
278 be helpful to constrain the HO<sub>x</sub> budget and its influence on daytime ozone.

### 279 3.6 Conclusion

280 SCIAMACHY was the first satellite instrument providing daily spectral solar irradi-  
281 ances (SSI) from the UV, visible and near infrared. The comparisons with other solar  
282 data from space and ground showed good agreement to within a few percent up to  
283 1700 nm (*Skupin et al.*, 2005a,b; *Pagaran et al.*, 2011a). Expressing SCIAMACHY  
284 irradiance variations over several solar rotations in terms of solar proxies for sunspot  
285 darkening and faculae brightening permits the extrapolation of SCIAMACHY SSI  
286 variations to the 11-year solar cycle scales (*Pagaran et al.*, 2009, 2011b). It was  
287 shown that about half of the 0.1% change in the solar constant over solar cycle  
288 23 has originates from the visible and IR spectral region (*Pagaran et al.*, 2009). A  
289 particular challenge is the solar cycle variation estimate for the near UV (300-400  
290 nm), where recent SIM observations (*Harder et al.*, 2009) indicate changes during  
291 solar cycle 23 that are much higher than expected from indirect SCIAMACHY ob-  
292 servations and other empirical models (assuming solar surface magnetic activity as  
293 primary driver for SSI variations) as well as observations from other satellite data in  
294 earlier solar cycles cycles (*Pagaran et al.*, 2011b).

295 SCIAMACHY limb ozone vertical profiles from 2003 to 2008 were analyzed for  
296 signatures of the 27-day solar rotation. It was found that this signature is highly  
297 variable in time and that even under solar maximum condition this signal can vanish  
298 for several months (*Dikty et al.*, 2010b). On average the sensitivity above 30 km is a  
299 0.2% ozone change per percent change in the 205 nm solar flux (important for ozone  
300 production) near solar maximum, which is smaller than found in earlier studies and  
301 prior solar cycles.

302 Daytime variations in tropical mesospheric ozone yield changes of up to 60%  
303 from the daytime mean based upon SABER ozone data and peak anomalies are gen-  
304 erally higher during equinox. SABER results were compared for the first time with  
305 an output of a chemistry climate model, here the HAMMONIA model (*Dikty et al.*,  
306 2010a). SABER ozone from the 9.6 $\mu$ m retrieval agrees qualitatively very well with  
307 HAMMONIA, however, little agreement was found between modeled and SABER  
308 temperatures above 0.01 hPa. The low temperature variations of a few degree during  
309 daytime may suggest that photochemical processes are the main driver for daytime  
310 ozone variations and to a lesser degree transport related to tides.

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