The Radiometer for Atmospheric Measurements (RAM)

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Overview

The Radiometer for Atmospheric Measurements (RAM) is operated by the Institute of Environmental Physics at the University of Bremen in cooperation with the Alfred-Wegener-Institute for Polar and Marine Research. It is part of the Network for the Detection of Stratospheric Changes (NDSC).

The Instrument is located at the Koldev vey-Statior Ny-Ålesund, Spitsbergen, at 78°55' N, 11°55' E. This observation site was chosen because of the special climatic conditions. The polar vortex is mostly located above Ny-Ålesund, which is essential for observing disturbed stratospheric chemistry.

The RAM is operational since November 1994. At present, it consists of three different radiometer frontends for measuring the ozone emission line at 142 GHz, the chlorine-monoxide line at 204 GHz and the water-vapour line at 22 GHz. One acousto optical spectrometer (AOS) is shared by the three frontends.

From the measured spectra, altitude profiles of the trace gas mixing-ratios are retrieved.



re 1: Inside the NDSC la r, in the cer Water-Vapour raciometer, in the center is the occure radiometer, and the intermeter is the intermeter in the chlorine-monoxide radiometer. The central rack carries the intermeter frequency chain and the acousto optical spectrometer. All these instrum observe the atmosphere through windows made of styropor.

Technical Data

	O ₃	CIO	H ₂ O
Receiver temperature	ambient	13 K	ambient
System noise temperature	3400 K	1150 K	200 K
Typical signal strength	15 K	0.05 K	0.2 K
Observation scheme	total power	reference	reference
Typical integration time	1 h	24 h	tbd
Single sideband filter	MPI	MPI	coaxial
Observation frequency	142.2 GHz	204.4 GHz	22.2 GHz
Spectral bandwith	1.65 GHz	0.96 GHz	0.96 GHz
Effective frequency resolution	1.6 MHz	1.6 MHz	1.6 MHz
Altitude range	1255 km	1635 km	tbd
Operational since	1993	1994	1999

Frontend

The idea behind microwave radiometry is the passive measurement of the rotational transition of thermal excited atmospheric trace gas molecules. The weak intensities of these signals require a special tech-nique for amplification and detection.

The basic concept of each radiometer is the same. The atmospheric signal is collected by a mirror and fed into a quasioptical system. Common elements of our quasioptics are a path length modulator for the extinction of standing waves, Martin-Puplett Interferometers as single-sideband-filters and cold and hot loads for calibration purposes. A mixer converts the hf signal down to the intermediate frequency of 8 GHz.

We use two different observation schemes. The total power method compares the atmospheric signal to cold and hot loads. The reference beam method compares two atmospheric signals at different observation angles. A compensating plexiglass sheet in one beam accounts for the longer path length through the atmosphere in the other beam.



Figure 2: The schematics of the ozone radiometer in total power geometry. The mirror switches be-tween the atmosphere, a cold and a hot calibration load. The Schottky diode mixer and the first HEMT amplifier stage are presently operated at ambient temperature.



Figure 3: The sch itics of the chl radiometer in reference beam geometry. In contra-to the ozone radiometer, the local oscillator optically coupled into the mixer. The whisk contacted diode mixer and the first HEMT amplifi stage are mounted inside a 13 K cryo system.



-Figure 5: Symbols of quasioptic and microwater vapour radiometer. operated in the reference b mode. In contrast to our othe diometers, a low noise preat fier is located in front of the mix neter. It is

Ζ $\overline{\langle}$ ellipsoidal mi

~~~ absorber

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X mixer

 $\bigcirc$ local oscillato

 $\triangleright$ amplifier

 $\approx$ 

band-pass filte

isolator

The rotating mirror wire grid

Backend

Our three radiometers share one backend, which consists of the intermediate frequency chain and an acousto optical spectrometer.

The purpose of the if-chain is the downcon-version from the first intermediate frequency of 8 GHz to the spectrometer input frequency of 2.1 GHz and the adjustment of the input power

About 20% of the observation time is used by the ozone radiometer, the remaining 80% are shared by the CIO and H O radiometers Although we measure throughout the year, CIO can only be measured in the arctic spring, because we subtract day- and night-time

measurements to eliminiate instrumental effects. Whenever the weather conditions are to bad for CIO, we measure  $H_2O$  instead.

A computer receives the data from the AOS and drives the attenuator, the mirror positioning units and the switches.

Our instrument is highly automated and requires only little technical attention.

Figure 5: The structure of our backend. A pair of switches selects the input signal. A portion of the signal is fed to a Powermeter. A comb signal generator can be coupled in for frequency calibration of the spectrometer. Once in a month, our data is stored on CD-ROM and shipped to Bremen for processing.





Figure 7: The schematics of an acousto optical spectrometer. The hf signal drives a piezo elec-trical transducer which generates a density grid in the connected deflector crystel. An expanded laser beam is defracted and the first order beam is projected on a CO-array. The deflection angle of the beam depends on the microwave-frequency and the intensity on the microwave-nover.

#### **Typical Results**



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rigure 8: Spectral measurement of the ozone rotational line at 142.175 GHz. March 20th 1997 at 17:00.

#### Figure 9: The retrieved ozone VMR profile and an ozone sonde Figure 10: Ozone volume-mixing-ratio profile timeseries from February 1997 to March 1998 over Ny-Ålesund.





Figure 11: A typical spectral measurement of the water-vapour line at 22 GHz. The blue line is a radiative transfer calculation of a given water-vapour distribution.



Figure 12: The red dots are a typical CIO mea surement with a superimposed baseline structur In comparison the blue line – a radiative transf calculation of a given CIO distribution.