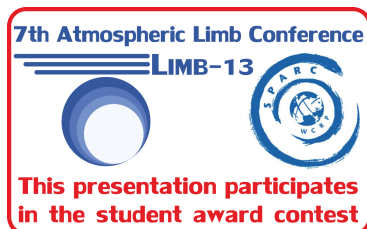


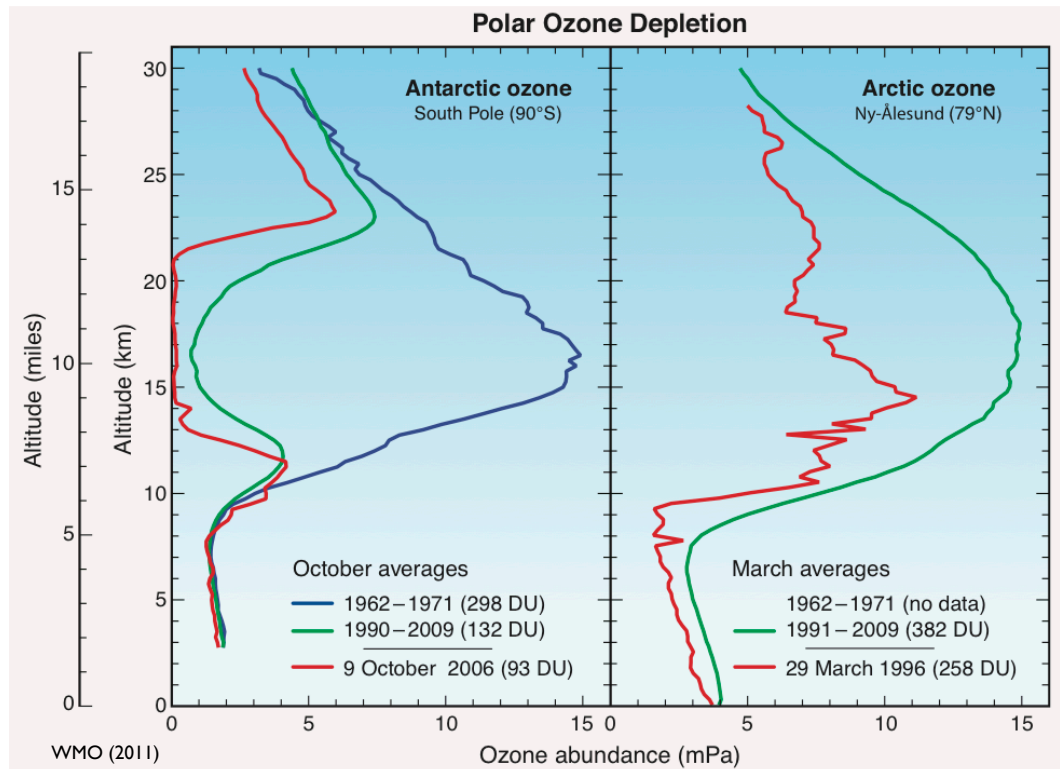
# Twelve years of Arctic ozone depletion observed by Odin/SMR

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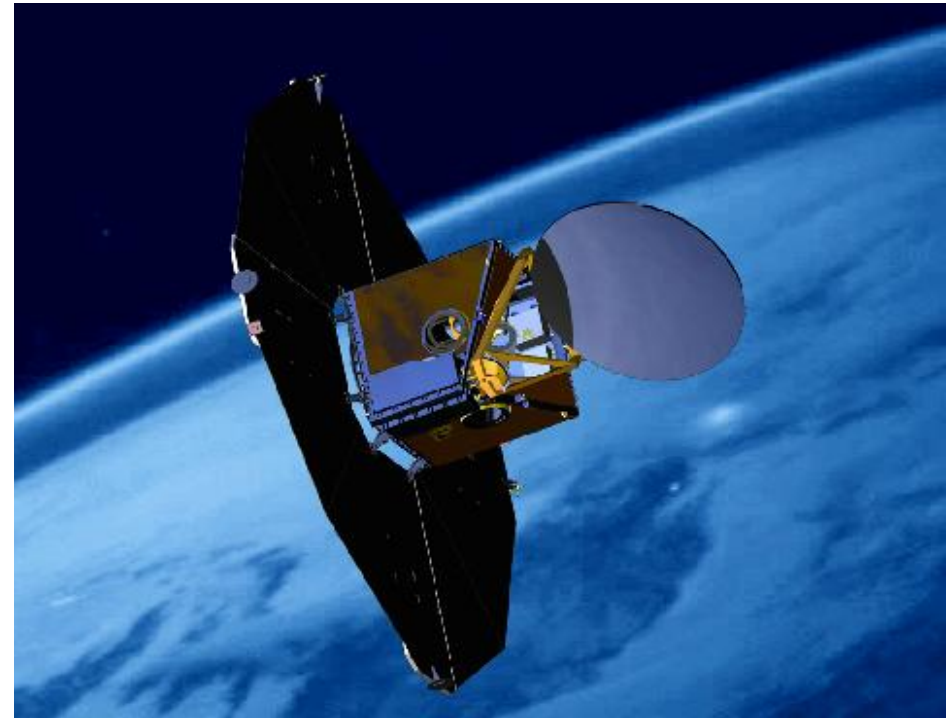
# Polar Ozone Depletion



- Since the early 80's, large scale ozone depletion has been observed in the lower stratosphere in the early spring. (Farman et al., 1985)
- Polar ozone loss is caused by catalytic species, such as Cl, NO<sub>x</sub> and Br
- The process of depletion is also related to the formation of PSC
- Huge Ozone loss such as Antarctic ozone hole was observed in 2011 winter (Manney et al., Nature, 2011)

# Odin/SMR

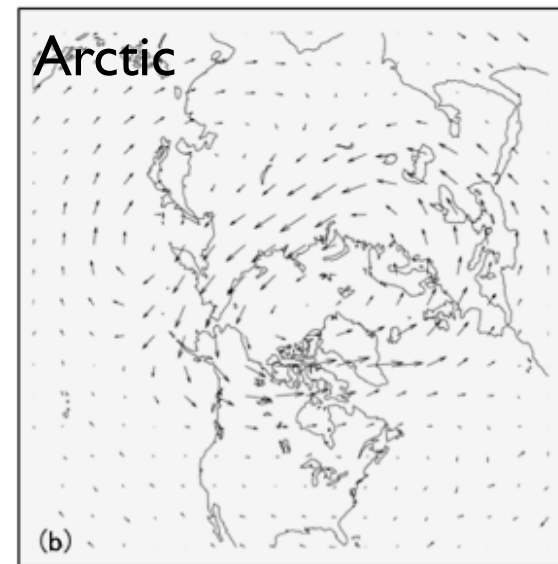
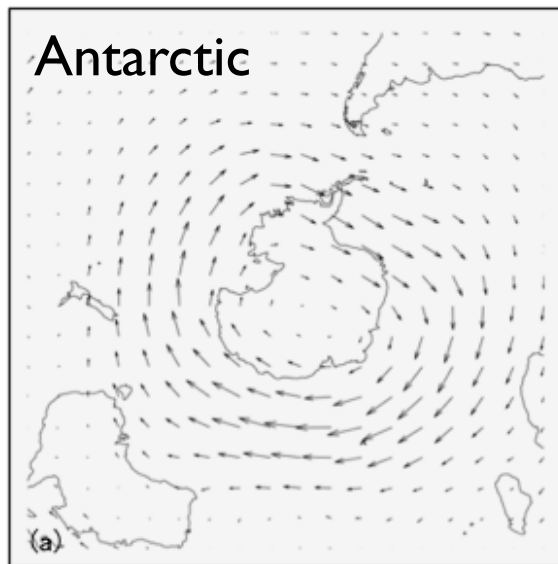
- Swedish-led satellite, collaborated with Canada, France and Finland
- OSIRIS (Optical Spectrograph and Infrared Imaging System)
- SMR (Sub-Millimetre Radiometer)
- launched on February 20, 2001



Long-term observation of stratospheric ozone can be used for trend study

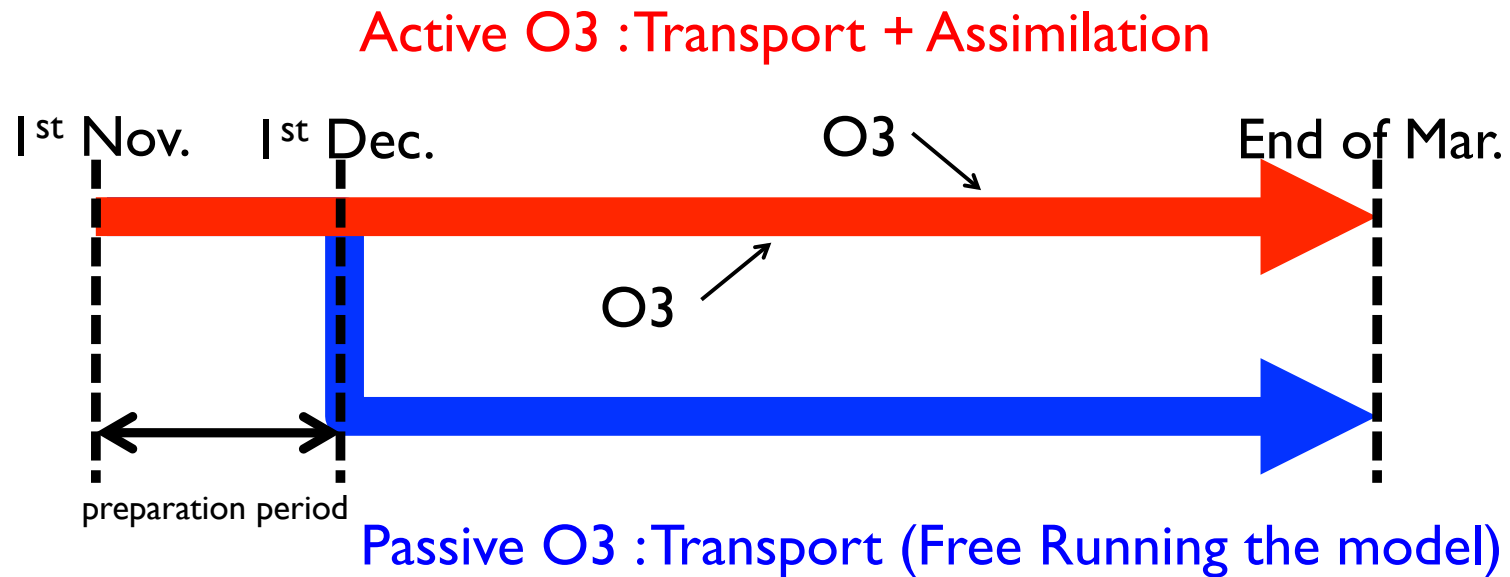
# Arctic Ozone depletion

- difficult to quantify the ozone depletion over arctic region due to instability of the polar vortex



Stratospheric wind patterns over the winter poles. (a) Antarctic winds on 2006-07-01. (b) Arctic winds on 2007-01-01. The wind data have been obtained from the European Centre for Medium-range Weather Forecasts, ECMWF [John, 2007]

# Methodology



$$O_{3loss} = O_{3active} - O_{3passive}$$

Possible to separate into “transport” and “chemical process” on Ozone VMR variation

# DIAMOND model

## (Dynamical Isentropic Assimilation Model for Odin Data)

- designed to simulate ozone horizontal transport in lower stratosphere (Rösevall et al., 2008)
- Square grid dedicated for the poles
  - Horizontal Advection
    - Transport on isentropic surfaces
    - Prather transport scheme (Prather, 1986)
    - ECMWF operational winds
  - Vertical Transport
    - 1st order upwind scheme
    - vertical motions calculated using SLIMCAT diabatic heating rate
  - Assimilation
    - Statistical interpolation
    - Modified Kalman Filtering
    - Transported analysis errors

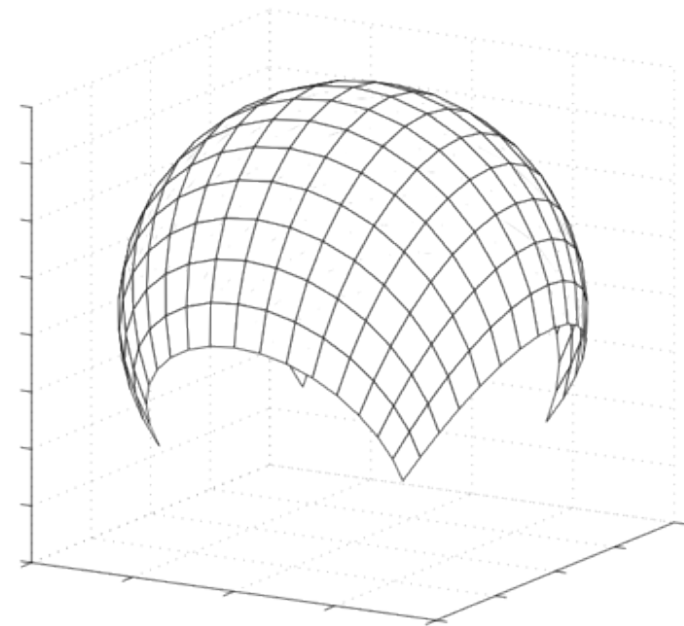
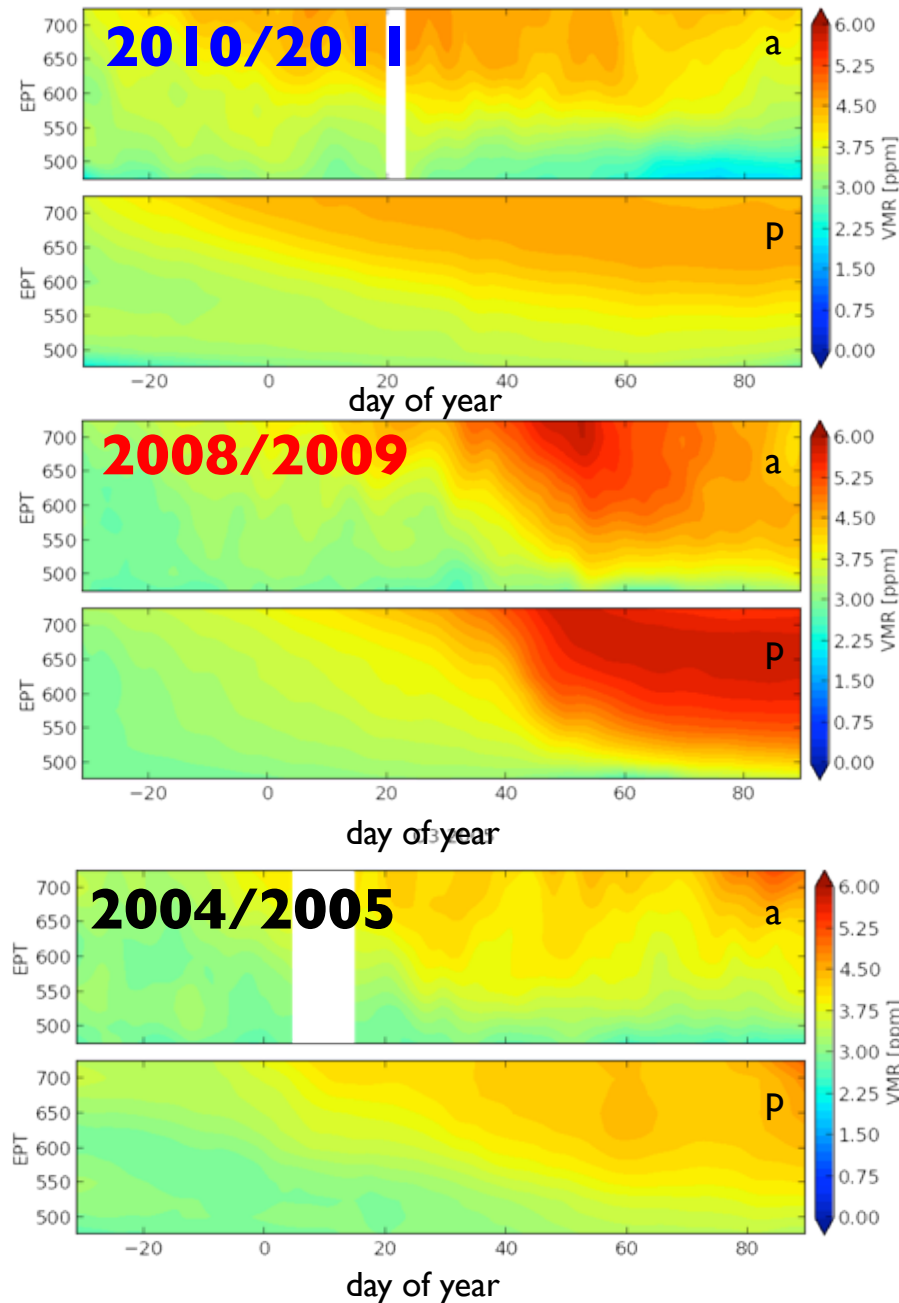


Fig. 2. The shape of the grid used in the DIAMOND model.

# Ozone data

- Odin/SMR
  - Retrieved from 501.5 GHz (Chalmers Retrieval Scheme v2.1, Urban et al., 2005)
  - Altitude range and resolution : 18 ~ 50 km, 2.5-3.5 km
  - Systematic Error : < 0.3 ppmv

# Vortex Mean of Ozone

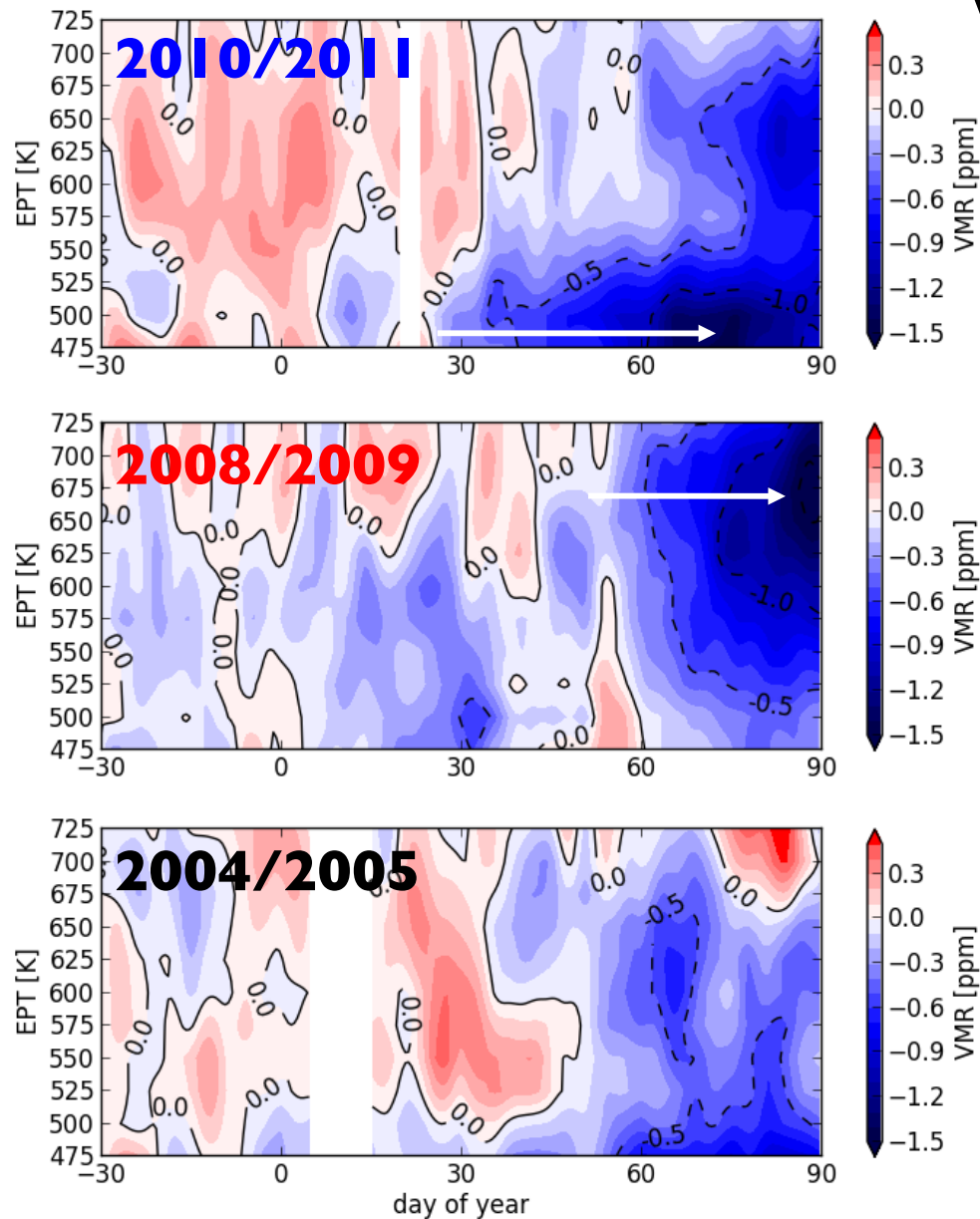


- Diabatic descent inside vortex is visible
- All ozone distributions can be categorized in 3 types

- 1) Cold Winter
- 2) Warm Winter
- 3) intermediate

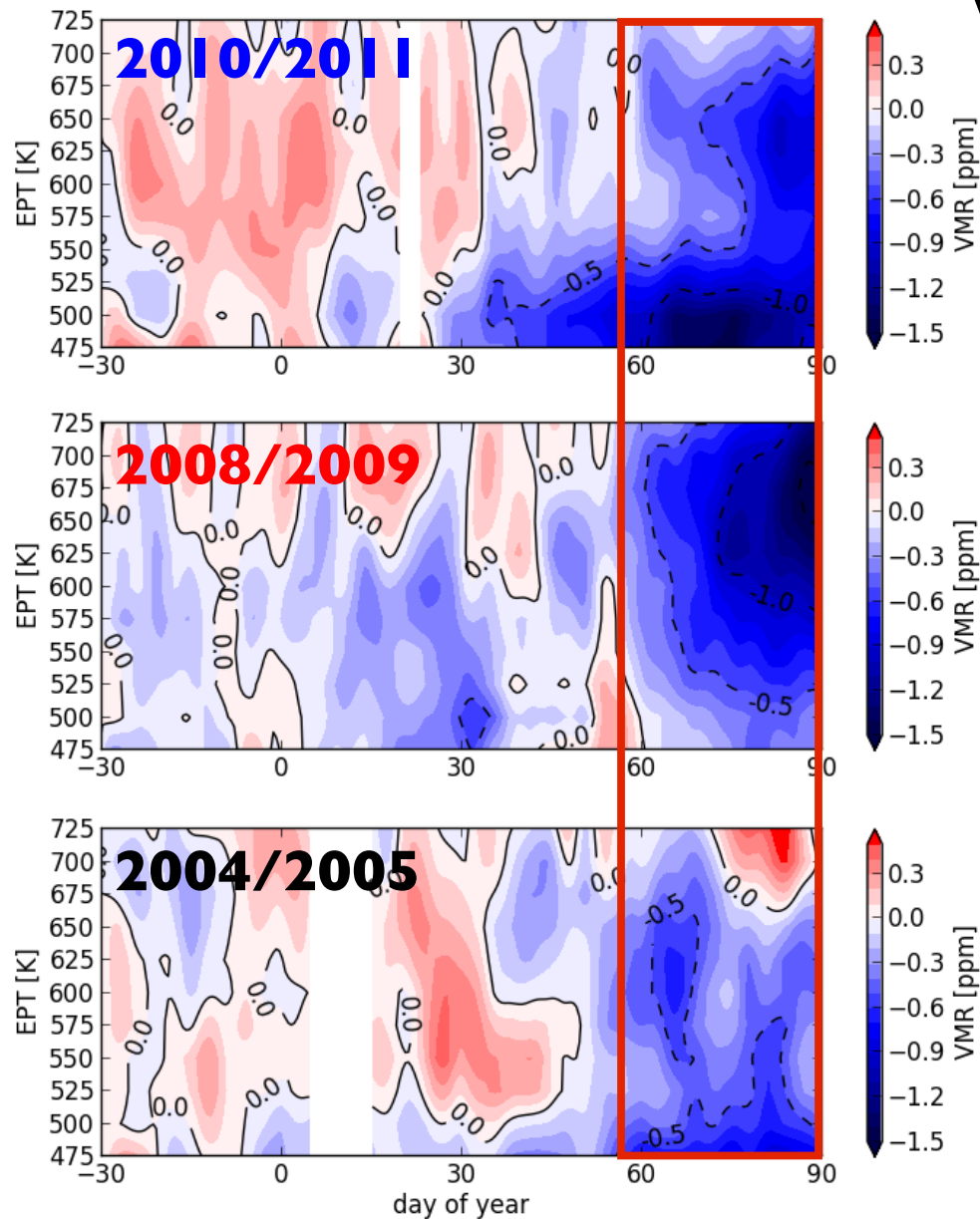


# Vortex Mean of loss



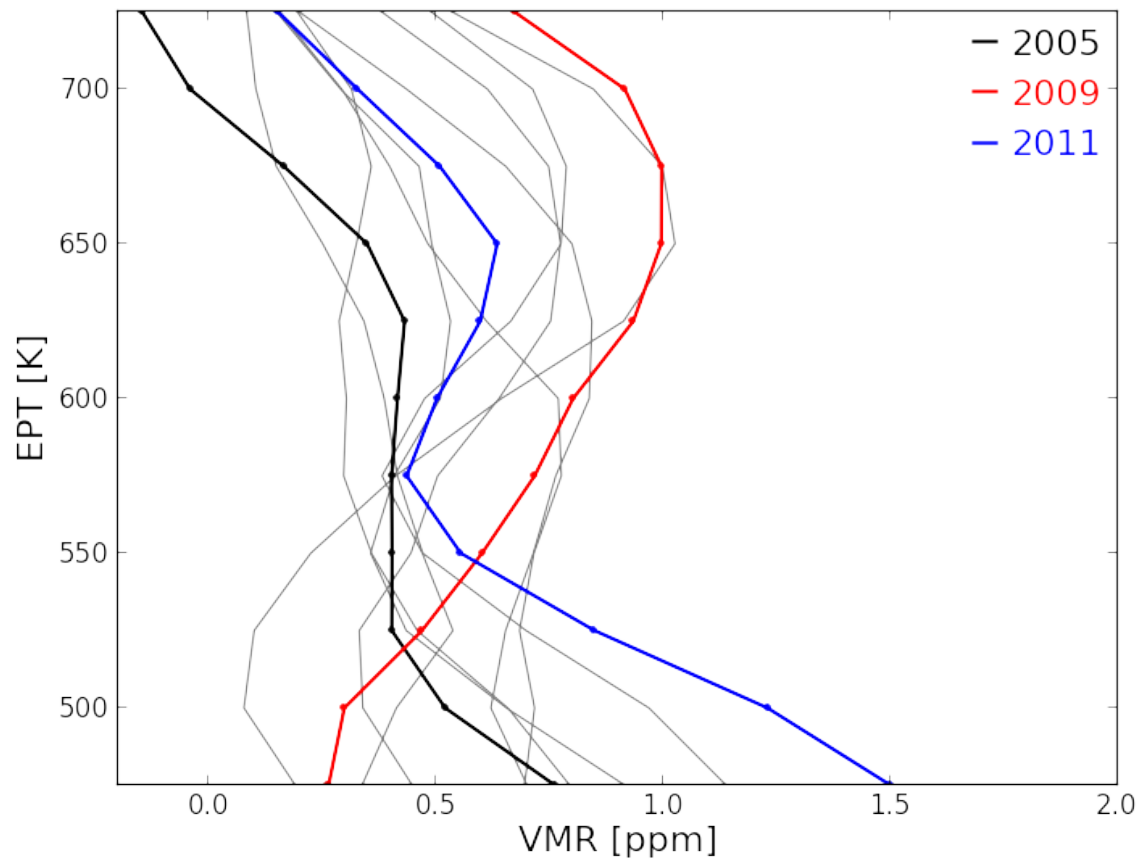
- Cold Winter
  - Ozone loss starts from beginning of February below 500K
- Warm Winter
  - major loss starts from beginning of March
  - around 650K

# Vortex Mean of loss

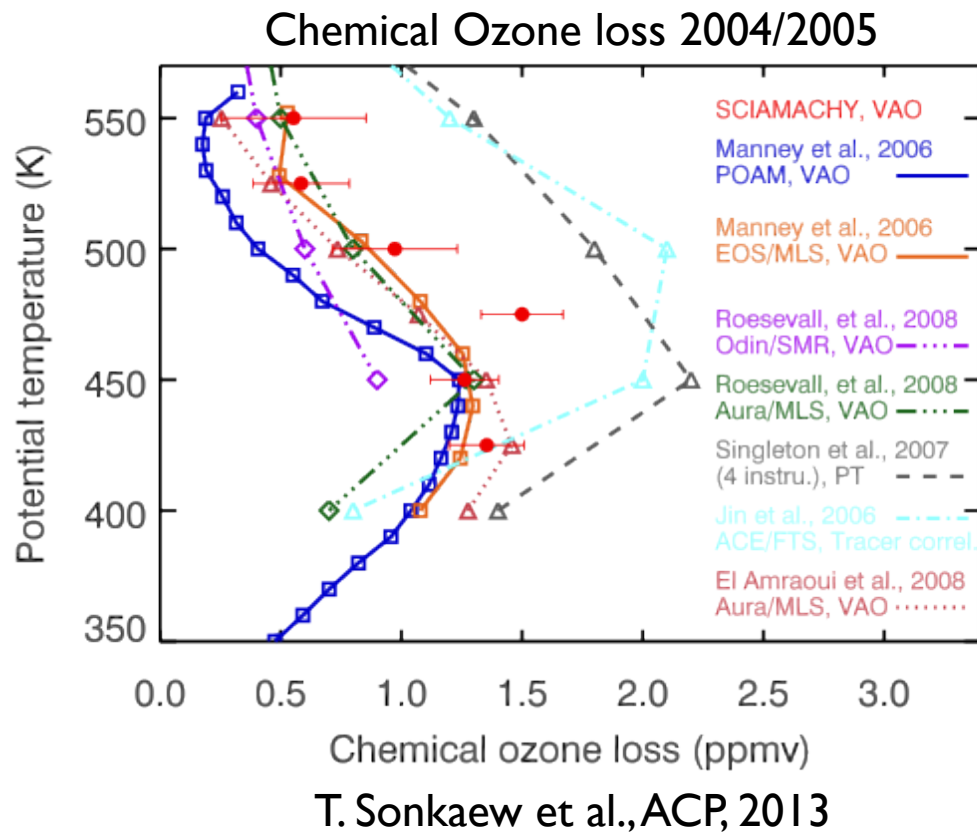


- **Cold Winter**
  - Ozone loss starts from beginning of February below 500K
- **Warm Winter**
  - major loss starts from beginning of March
  - around 650K

# Monthly Average of Chemical ozone depletion for March

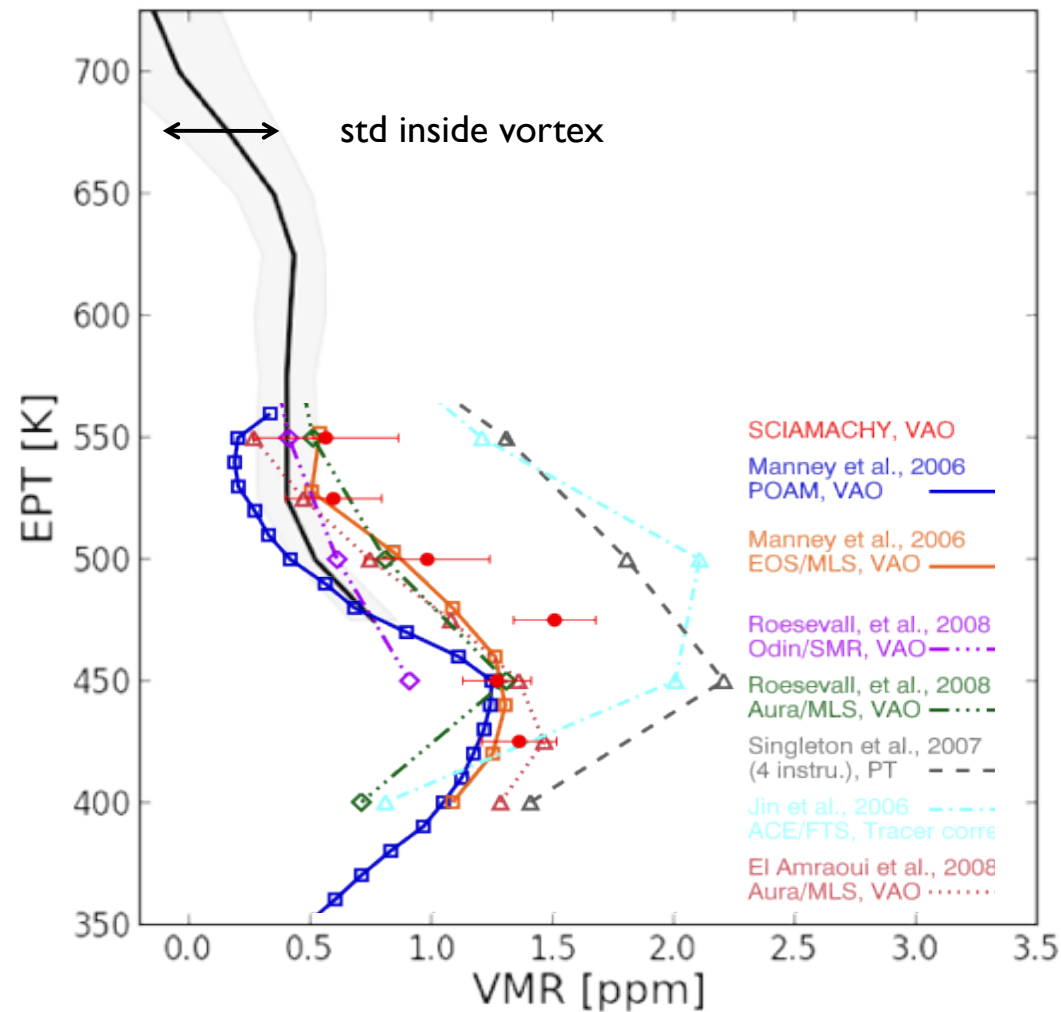


# Comparison with other Instruments and Methods



Ref	Instrument	Technique
<b>T. Sonkaew et al., 2013</b>	<b>SCIAMACHY</b>	<b>Vortex Average</b>
<b>Manney et al., 2006</b>	<b>POAM</b>	<b>Vortex Average</b>
<b>Manney et al., 2006</b>	<b>EOS/MLS</b>	<b>Vortex Average</b>
<b>Rösevall et al., 2008</b>	<b>Odin/SMR</b>	<b>Original DIAMOND</b>
<b>Rösevall et al., 2008</b>	<b>Aura/MLS</b>	<b>Original DIAMOND</b>
<b>Singleton et al., 2007</b>	<b>POAM III, SAGE III, EOS MLS, ACE-FTS, and MAESTRO</b>	<b>Passive Subtraction</b>
<b>Jin et al., 2006</b>	<b>ACE/FTS</b>	<b>4 different techniques</b>
<b>El Amraoui et al., 2008</b>	<b>Aura/MLS</b>	<b>Vortex Average</b>

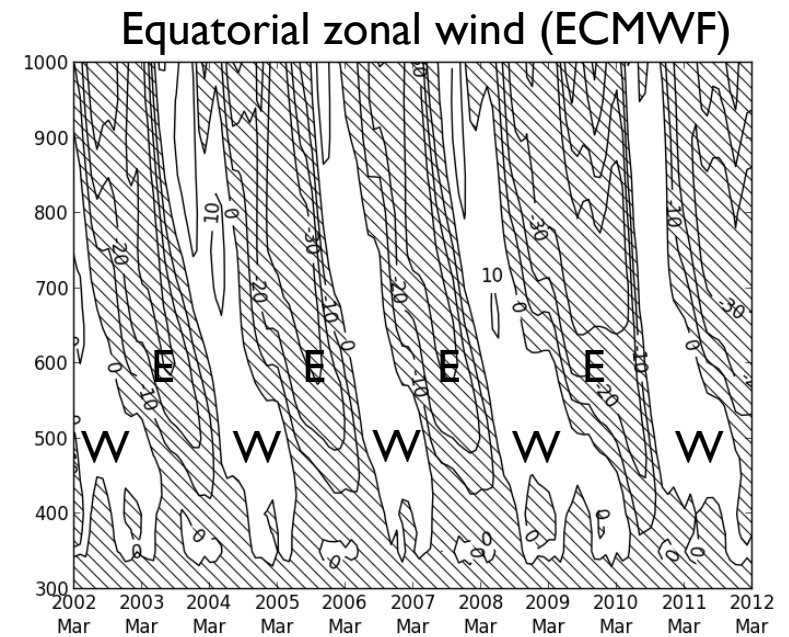
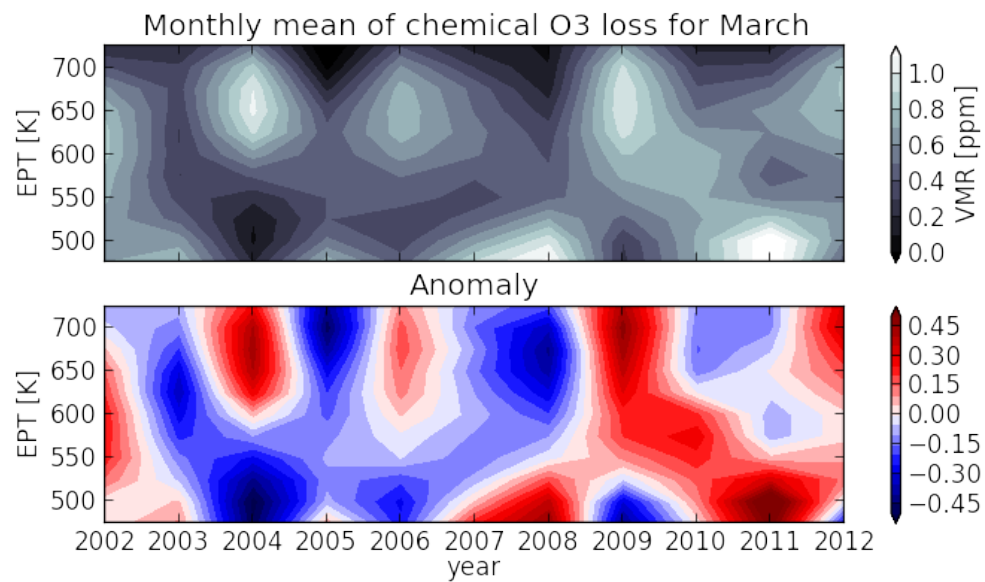
# Comparison with other Instruments and Methods



## Under estimate the loss

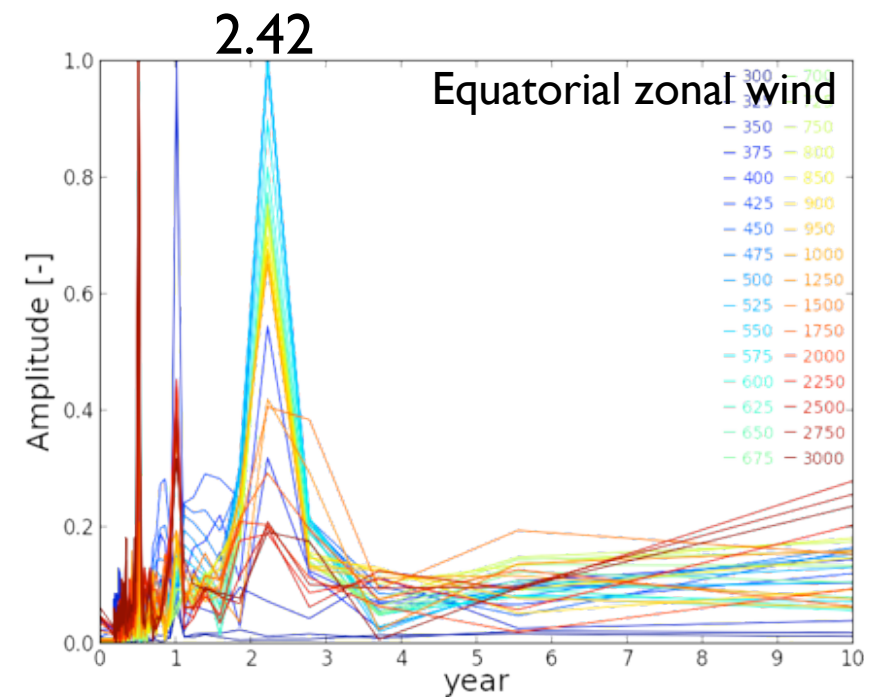
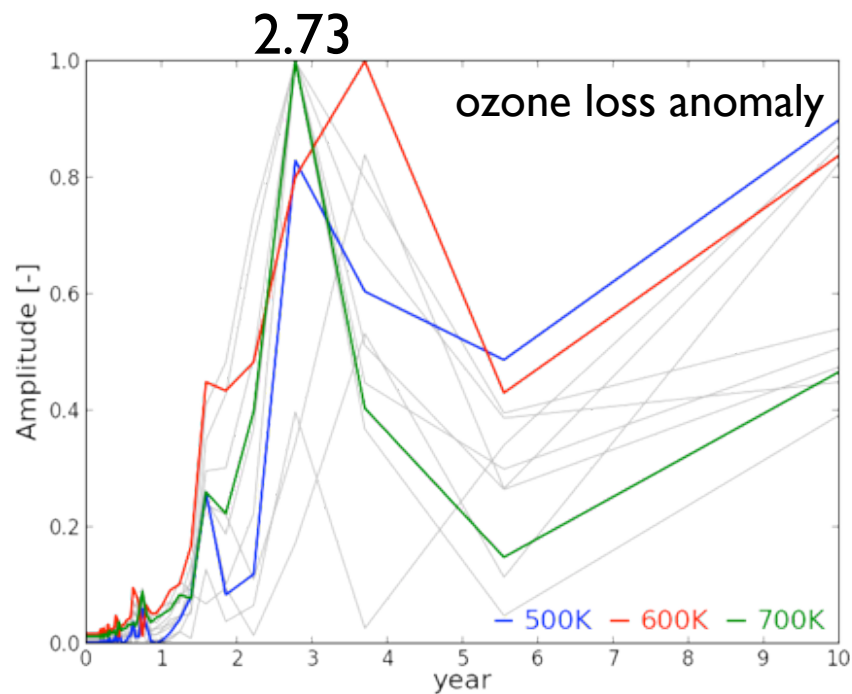
- Odin/SMR Ozone biased?
- Due to the under estimate of diabatic descent inside vortex?

# Variation of loss



Periodic pattern is similar to QBO in equator

# Spectrum of Ozone loss anomaly



# Summary

- Odin/SMR is observing stratospheric ozone from 2001
- Arctic chemical ozone losses of Northern winters from 2002 to 2012 were quantified by using DIAMOND model
- Comparison of derived ozone loss in 2004/2005 winter agreed with the other methods in error range
- Losses from our method for all winters are lower than other instruments
- The peak of ozone loss is shown below 500K in cold winter, while it appears higher in warm winter
- Variation of ozone depletions of each winters seems to have similar pattern with QBO



# Appendix

# How to combine model and measurement ?

- Assimilation is usually used for meteorological prediction
- Measurement ingested into numerical model to estimate the optimal state
- Estimated ozone VMR and assimilation error are given as

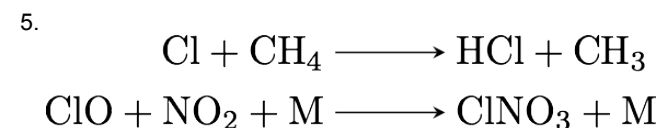
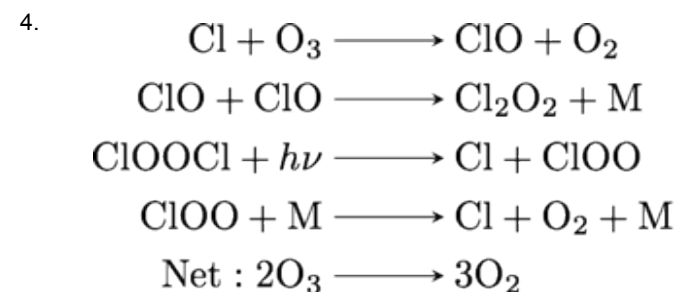
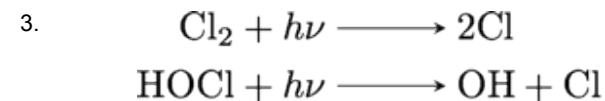
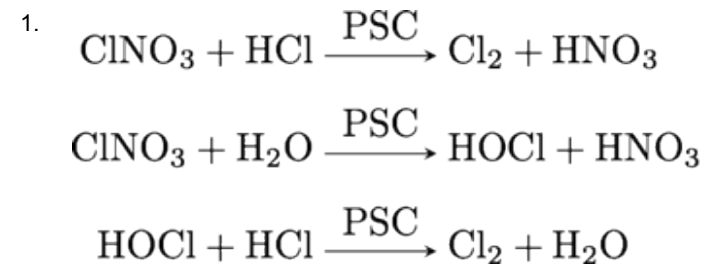
$$\hat{x} = \left( \frac{1}{\sigma_{model}^2} + \frac{1}{\sigma_{meas}^2} \right)^{-1} \left( \frac{x_{model}}{\sigma_{model}^2} + \frac{x_{meas}}{\sigma_{meas}^2} \right)$$
$$\sigma = \left( \frac{1}{\sigma_{model}^2} + \frac{1}{\sigma_{meas}^2} \right)^{-1}$$

- Linear growth of error with time can be used to approximate the uncertainties due to the imperfection of transport model and chemical processes in Stratosphere

$$\sigma(t + \Delta t) = \sigma(t) + k_{\sigma} \cdot \Delta t$$

# Chemical process of Ozone depletion

- (1) Cl<sub>2</sub> and HOCl are released by heterogeneous reactions on the surface of PSCs.
- (2) Cold winter (<-80oC) required to form PSCs
- (3) In early spring, Cl and ClO are produced by the photo-dissociation of Cl<sub>2</sub> and HOCl
- (4) Cl catalytic destruction cycle
- (5) The cycle stops by reactions to CH<sub>4</sub> or NO<sub>2</sub>



# Vertical transport

$$\psi(z, t + \Delta t) = \psi(z, t) \left(1 - \omega \frac{dt}{dz}\right) + \psi(z - \Delta z, t) \omega \frac{dt}{dz}$$

