

Limb cloud flagging

Sciamachy CLOud Detection Algorithm (SCODA)

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Contents:

- **Summary of Approach**
- **Status/Input for processing**

Limb Cloud Flagging

Color index (CI) approach:

Difference in relative contributions of scattering by air molecules (*Rayleigh scattering*) and the cloud particles (*Mie scattering*) at two different wavelengths

Considerations:

- Radiation < 400 nm not used (atmosphere optically thick for scattering in the upper troposphere)
- Windows with strong molecular absorptions not used
- Radiation \sim NIR used to distinguish the thermodynamic phase (spectral dependence of absorption differ appreciably between ice and water ~ 1670 nm)

SCODA Algorithm:

1- Color index profile (CI):

Ratios of spectrally integrated intensities in a defined wavelength band

$$R_c(TH) = \frac{I(\lambda_1, TH)}{I(\lambda_2, TH)}$$

Sensitivity to cloud determination

2- Color index ratio (CIR):

Starting from the lowest TH, pairs of two adjacent color indices divided, maximum shows the cloud top height (CTH)

$$\Theta(TH) = \frac{R_c(TH)}{R_c(TH + \Delta TH)}$$

$\Delta TH = 3.3 \text{ km}$

SCODA is part of operational retrieval chain since SCIAMACHY data Lv2 V5.04 (support to operational limb trace gas retrieval)

SCODA Clouds Types:

- Tropospheric (Water & Ice)
- Polar stratospheric (PSCs)
- Noctilucent (NLCs, North)

Wavelength (bands) ratio:

Noctilucent clouds:

265nm & 291nm

Cloud detection (water):

750nm & 1090nm

Phase determination (ice):

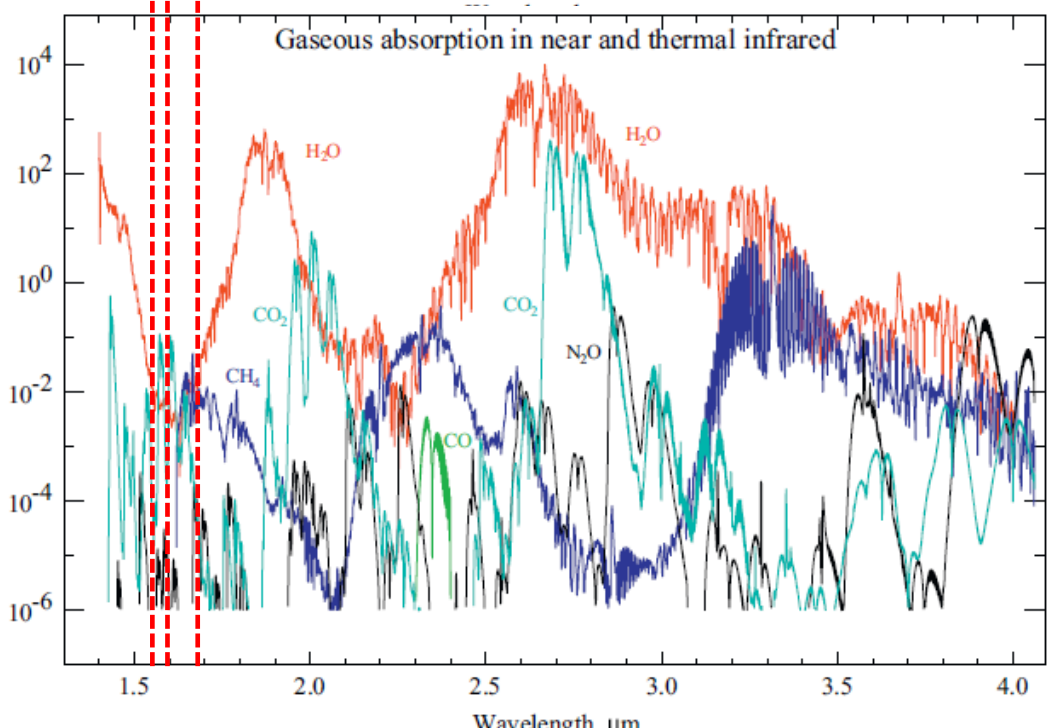
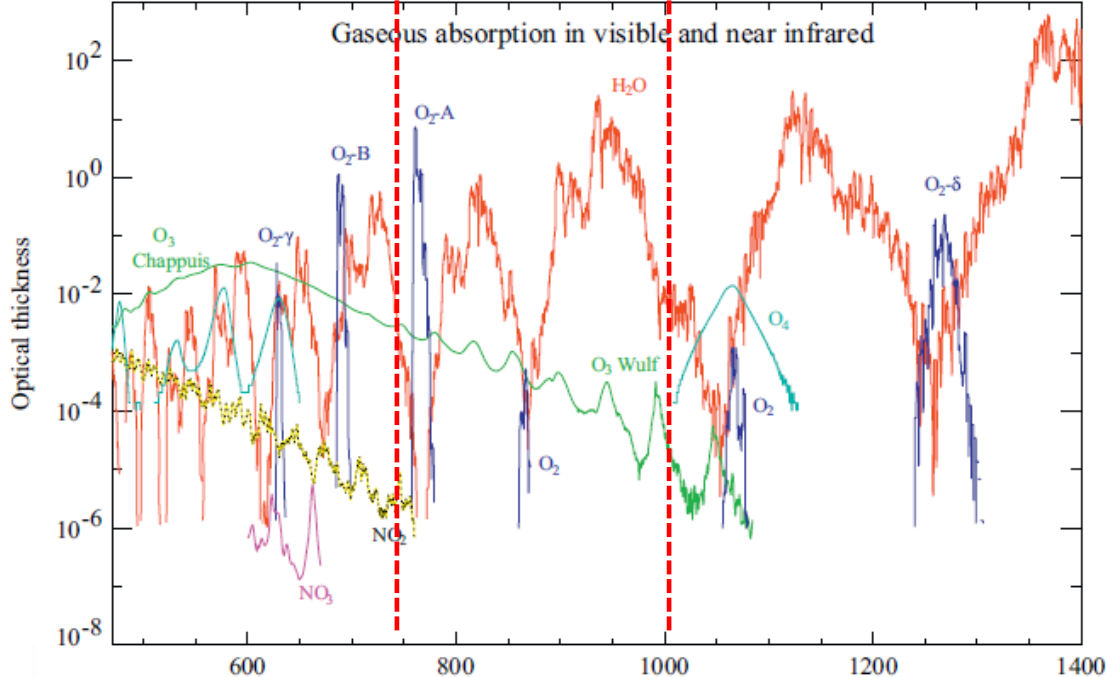
1550nm & 1685nm (IUP)

1550nm & 1630nm (Operational)

In addition limb water vapour retrieval at IUP uses 1550/1090 nm for cloud detection.



Rozanov V. V. et al., jqsrt 2014



SCODA flags (Operational processor):

Cloud type	CIR	Threshold	Flag	Description
Normal (water)	1090nm/750nm	$CIR < 1.4$	0	No clouds
		$1.4 < CIR < 2$	1	Partial
		$CIR > 2$	2	Full
		$CIR >>$	3	Bad data or >> Max. TH
PSCs	1090nm/750nm	$CIR < 1.3, lat. > 50^\circ, 15 < TH < 30$	0	No PSCs
		$CIR > 1.3, lat. > 50^\circ, 15 < TH < 30$	1	PSCs
**Ice	1630nm/1550nm	CIR?	0	Water cloud
		$CIR > ??$	1	Ice cloud
		$TH >> \text{ or } ??$	2	Bad data (>> Max.TH)

Limb Cloud flagging optimization:

➤ Improvement in detection sensitivity:

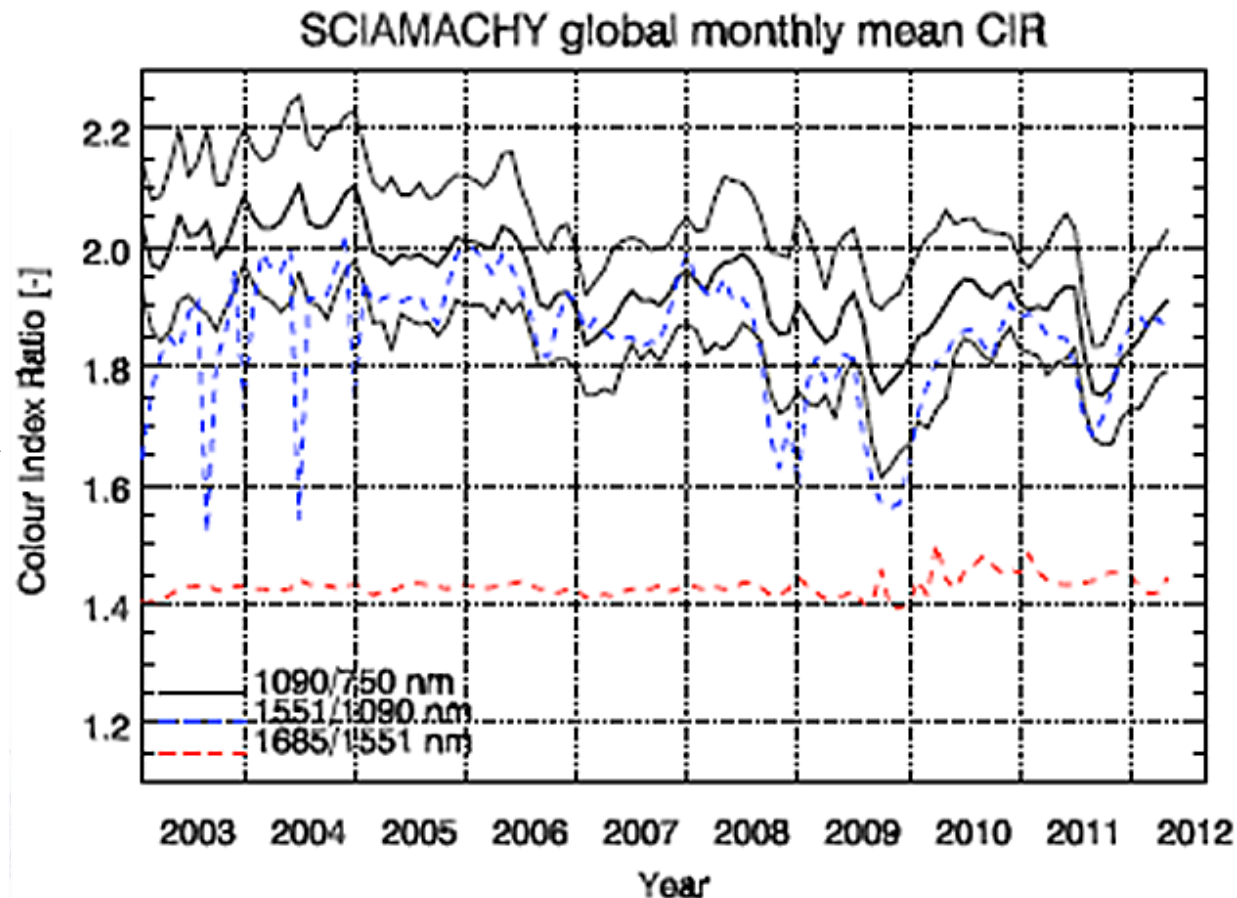
Optimize phase detection, issues as S/N decreasing with altitude, bad pixels in channel 6+

➤ Distinguish aerosols from clouds: CIR 1090/750 misidentifies high aerosol layers as clouds, CIR ~1670/1550 almost not effected

2008, 2009 and 2011
→ Volcanic eruptions

2003, 2004 and 2008
→ Decontamination phases

Eichmann K. U. et al 2015, submitted amt



Clouds discrimination Study

CI instead of CIR:

- Ratio of reflectance 1550/1670 (here normalized by Rayleigh reflectance, albedo=0), ratios improve on S/N

Note: 1670 +- 20 nm window is used, bad and RTS pixels flagged in real data

- At a given TH if ratio < 0.7, cloud present at this TH or above. All THs between ground and 20 km scanned

Flag:

- All events with cloud detected between 0-20 km
 - Consecutively ignore lowest THs until 15 km, get all event with clouds between TH_i-20 km
- V7 used with calibration setting of V8, bad pixel mask from L1b is implemented

SCIATRAN Simulations (1550/1670, Rayleigh normalized)

➤ 1 set with varying albedo/boundary layer/trop./stratospheric aerosol

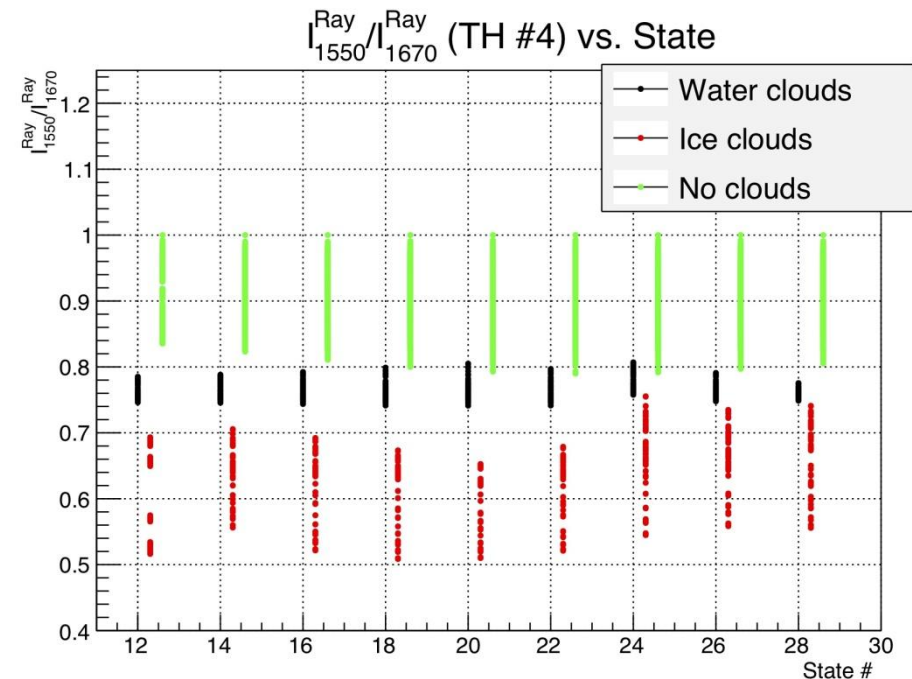
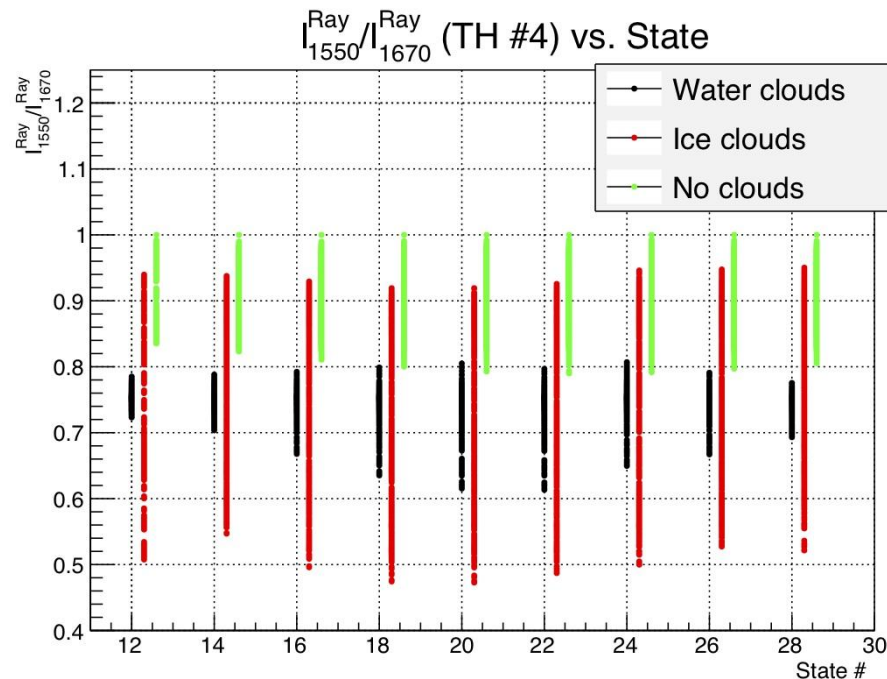
➤ 1 set with clouds at different altitudes in addition

Water clouds: 5-7, 9-11, 11-13, 13-15 km, OD=0.5-50

Ice clouds: 13-15, 15-17, 17-19, OD=0.001-0.5

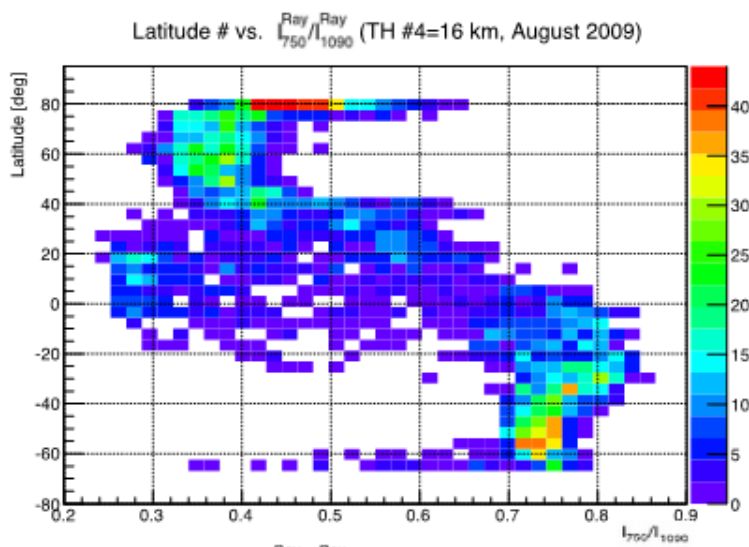
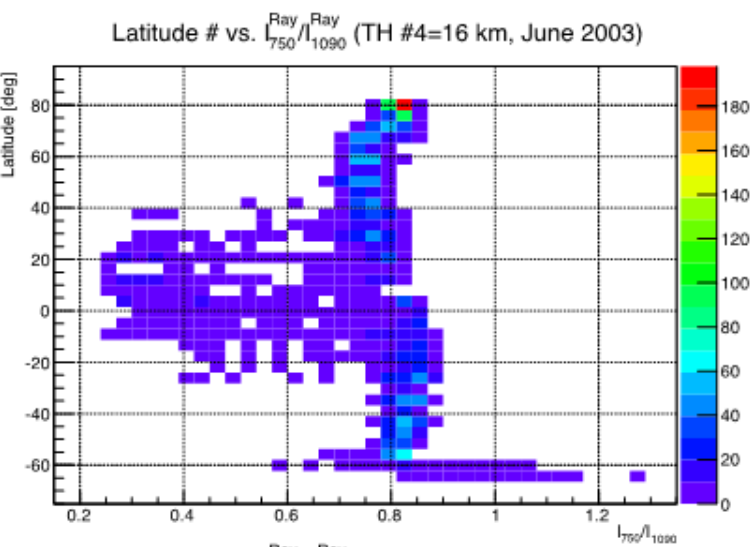
Each with 2 different size parameters

Discrimination is difficult unless similar background conditions and/or known OD

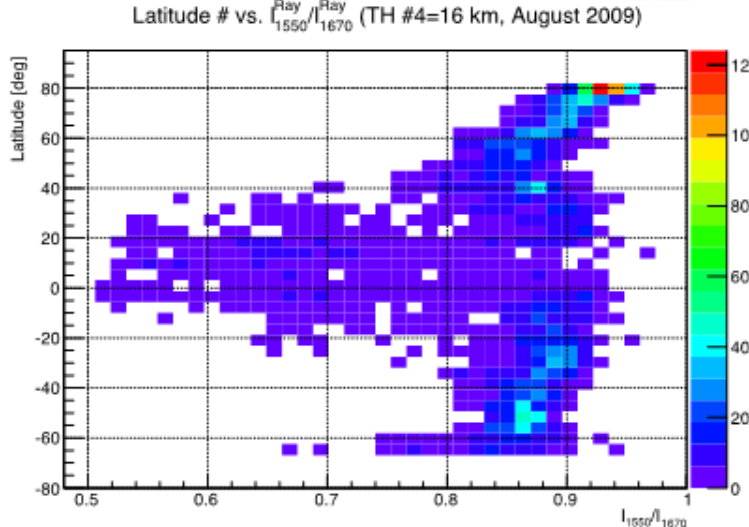
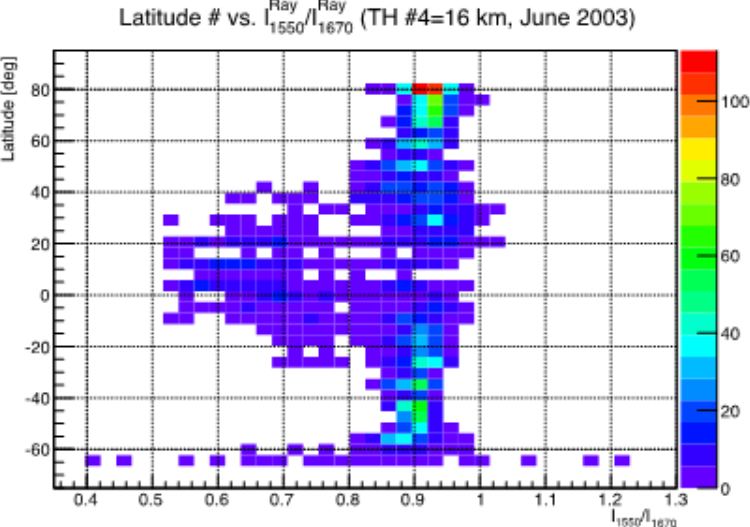


SCIAMACHY Limb data (ratio vs. lat.)

Rayleigh normalized ratio 750/1090 misidentifies high aerosol load as clouds, changes distribution (Sarychev eruption June, 2009), 1550/1670 unaffected



750/1090 nm

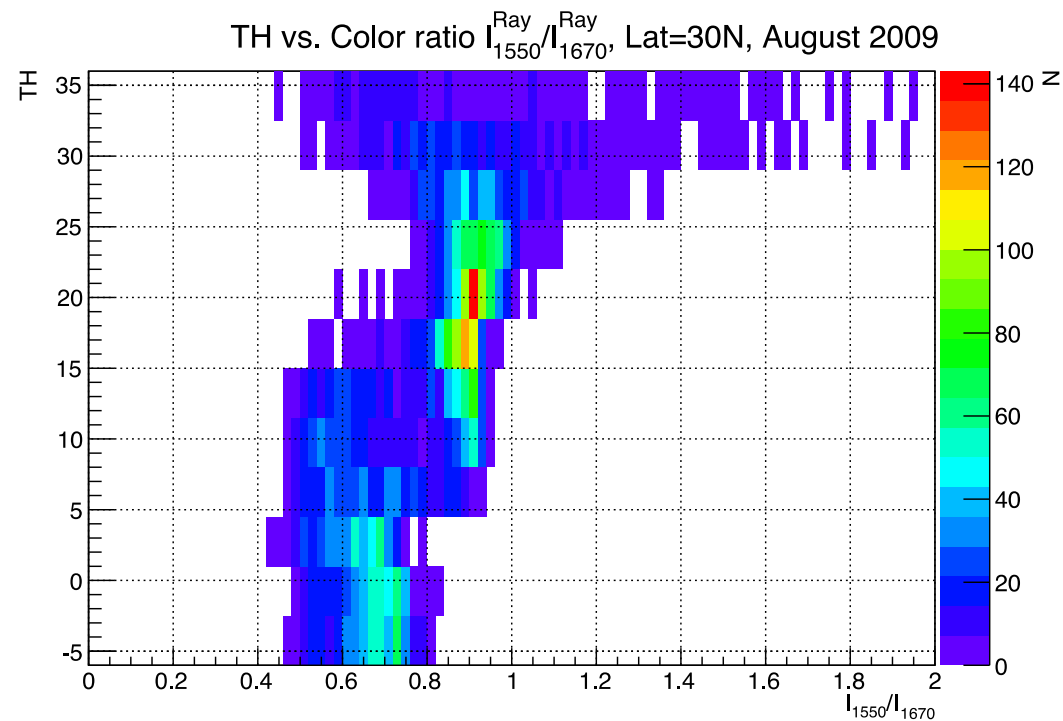
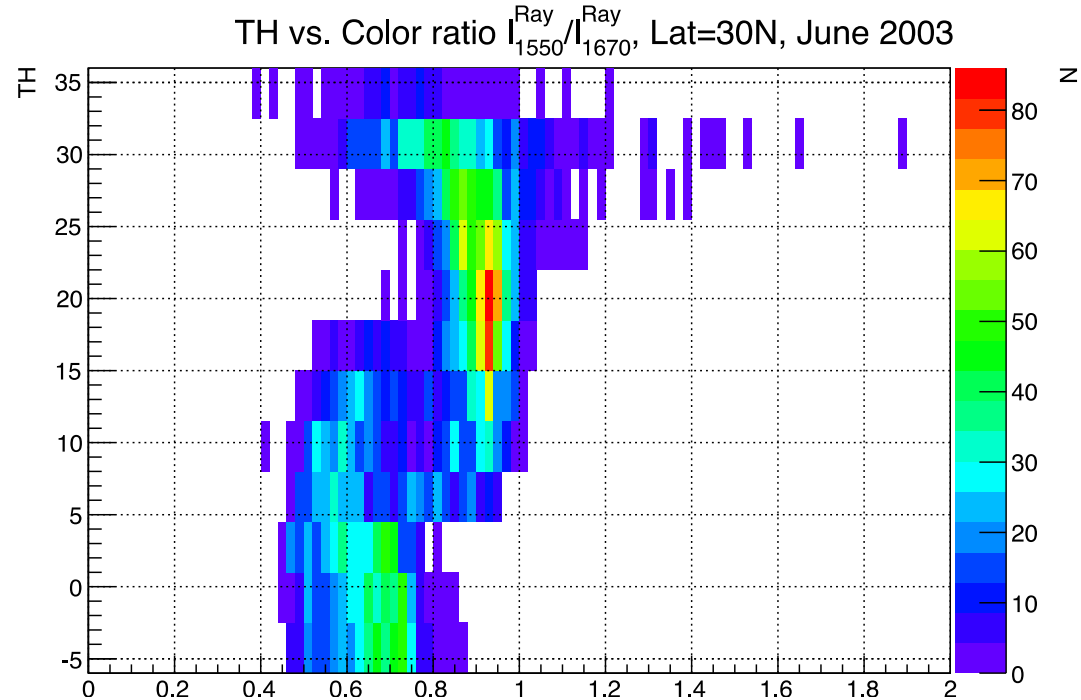


1550/1670 nm

SCIAMACHY Limb data (ratio 1550/1670 vs. TH)

No qualitative difference between the two years, although aerosol load is much higher in 2009.

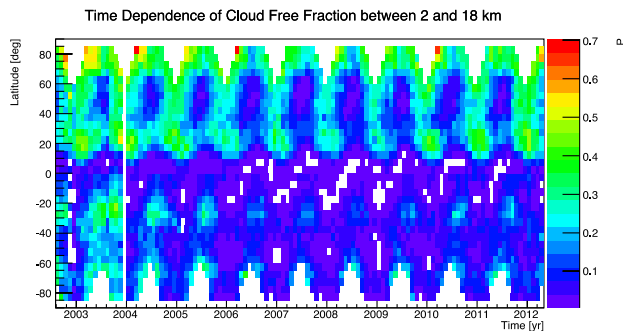
Differences from 30 km up.



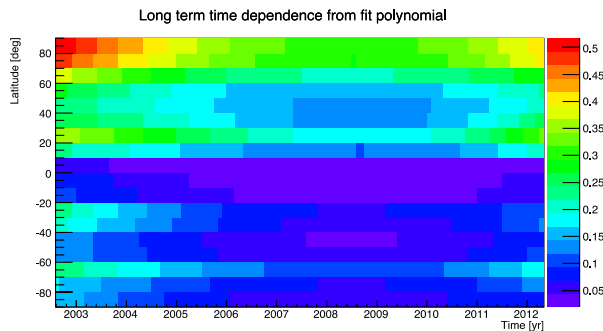
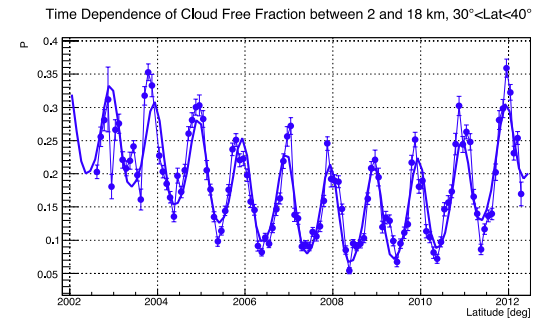
In addition

Cloud frequencies in Limb (monthly zonal means):

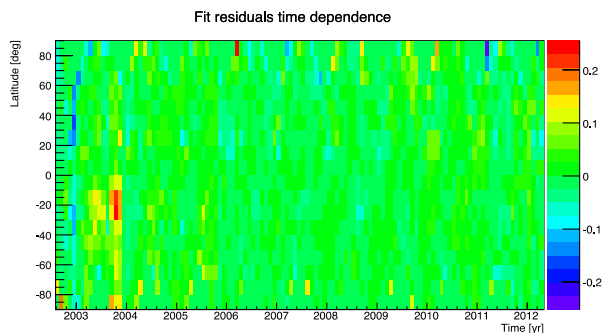
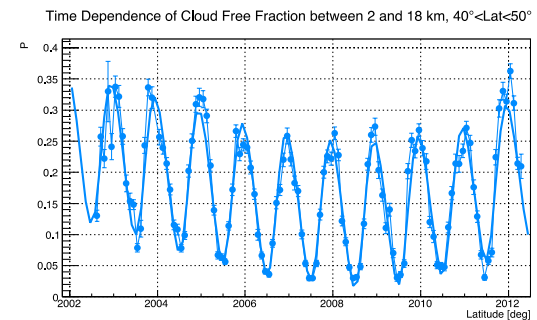
Fraction of cloud free profiles is strongly time/lat dependent



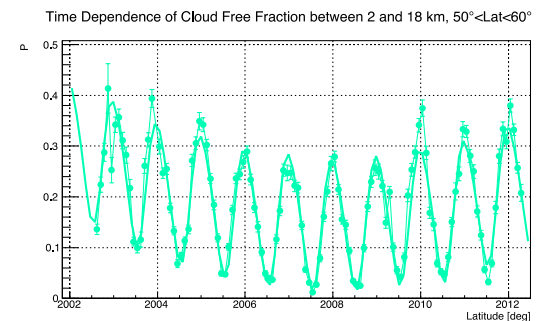
30N – 40N



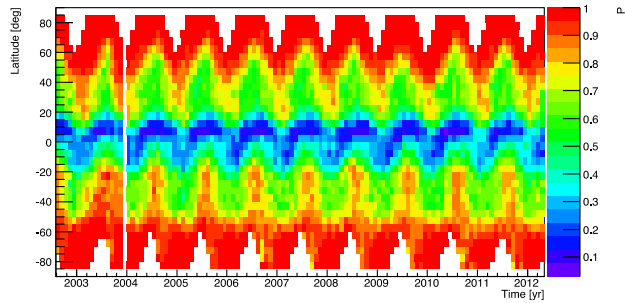
40N – 50N



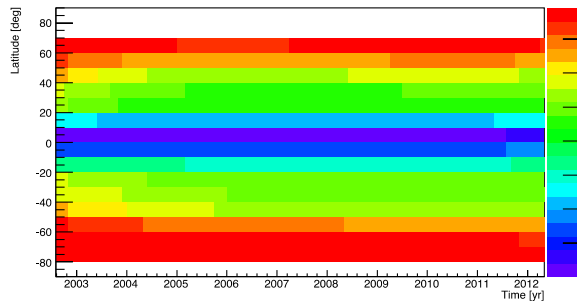
50N – 60N



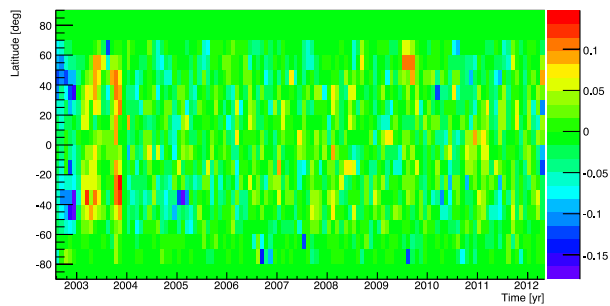
Time Dependence of Cloud Free Fraction between 12 and 18 km



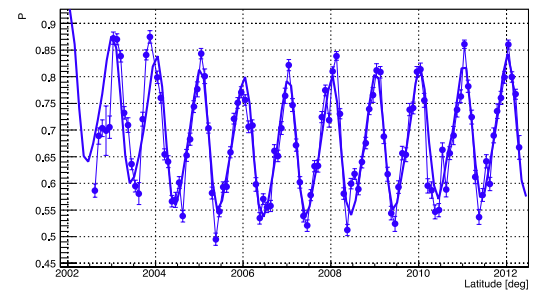
Long term time dependence from fit polynomial



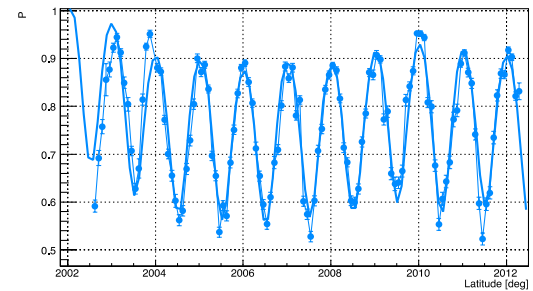
Fit residuals time dependence



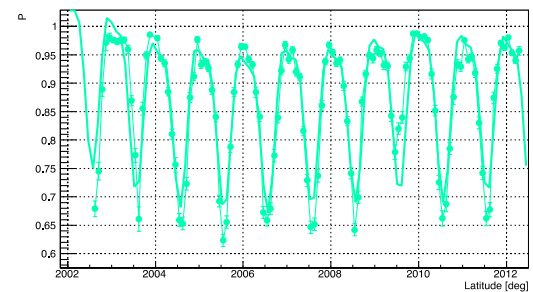
Time Dependence of Cloud Free Fraction between 12 and 18 km, $30^\circ < \text{Lat} < 40^\circ$



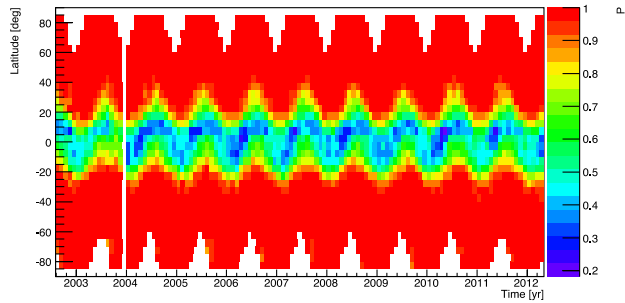
Time Dependence of Cloud Free Fraction between 12 and 18 km, $40^\circ < \text{Lat} < 50^\circ$



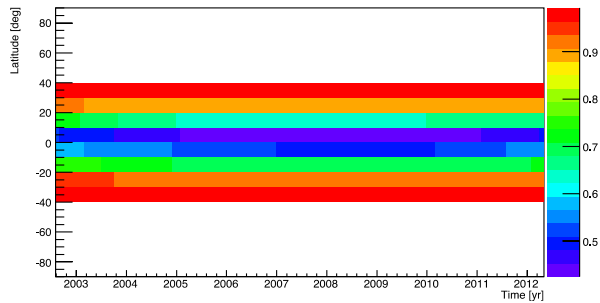
Time Dependence of Cloud Free Fraction between 12 and 18 km, $50^\circ < \text{Lat} < 60^\circ$



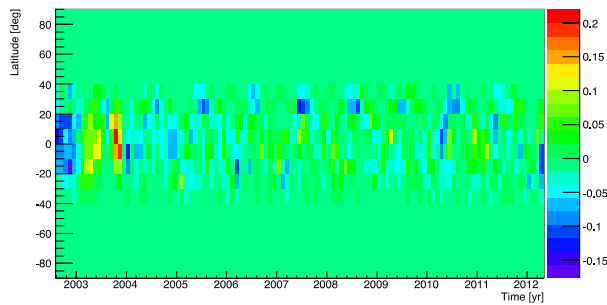
Time Dependence of Cloud Free Fraction between 15 and 18 km



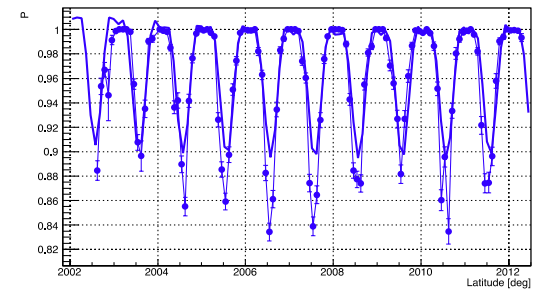
Long term time dependence from fit polynomial



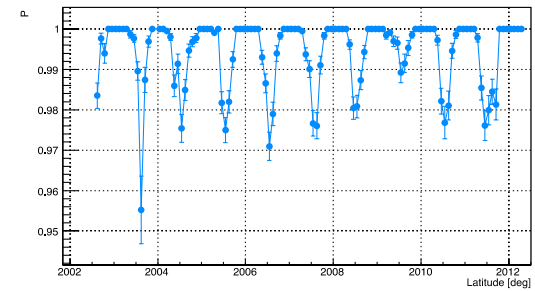
Fit residuals time dependence



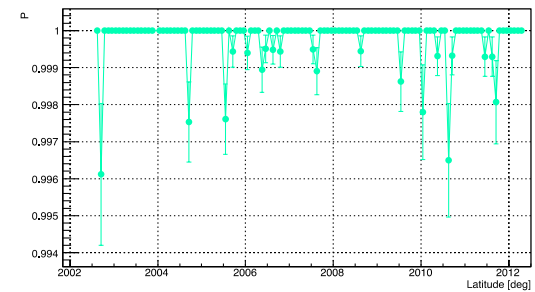
Time Dependence of Cloud Free Fraction between 15 and 18 km, $30^\circ < \text{Lat} < 40^\circ$



Time Dependence of Cloud Free Fraction between 15 and 18 km, $40^\circ < \text{Lat} < 50^\circ$



Time Dependence of Cloud Free Fraction between 15 and 18 km, $50^\circ < \text{Lat} < 60^\circ$



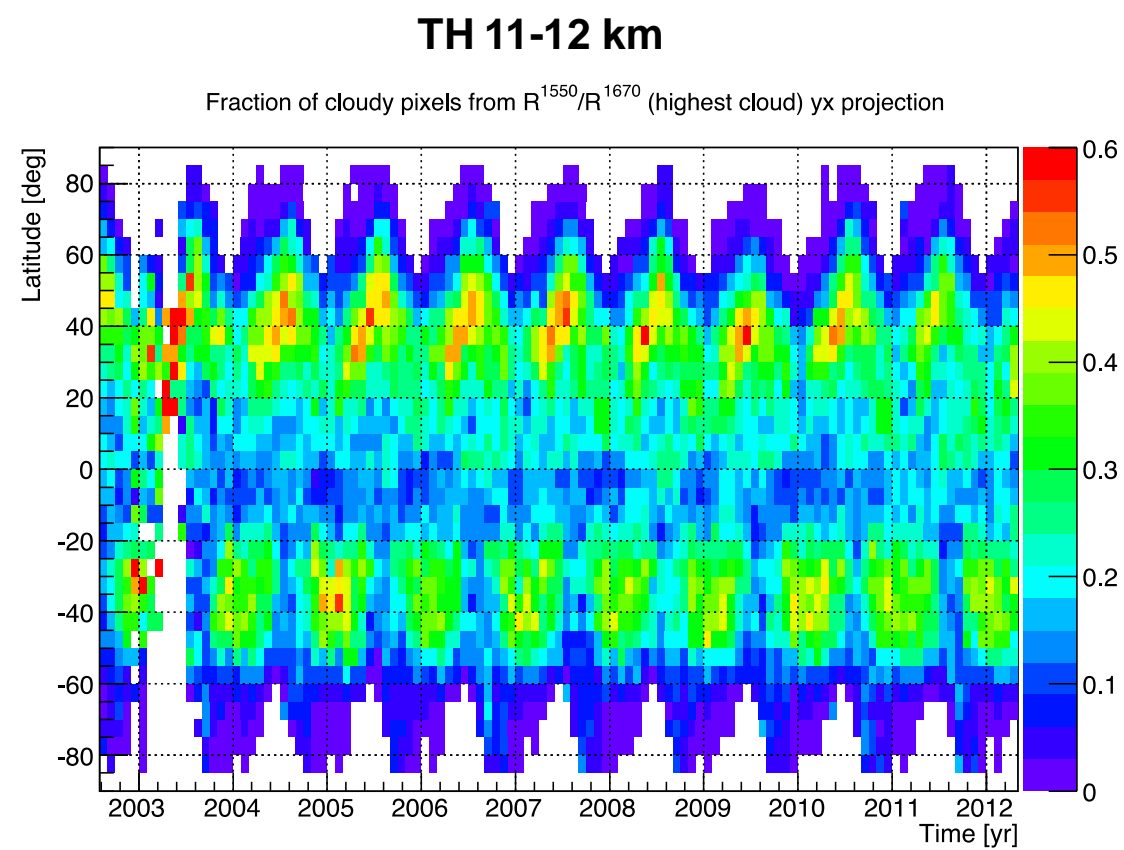
Several issues when analyzing time series:

- **Degradation/background:** wrong degradation or offset correction could shift the distribution (changing purity and efficiency of samples)
- **Tangent height sampling**
- **Noise** (gets worse if dark currents increase)
- Though 1550/1670 ratio is less sensitive to aerosols than the 750/1090, there is some sensitivity

Effect of different flagging:

only flag the highest TH where the cloud is detected, $R < 0.75$, cloudy fraction

Still strong time dependence



Possible hick ups:
decontaminations 2002-2004

TH sampling variations in 2002/2003

Orbit change (TH scanning regime)

Volcanic eruptions

Latitude dependence may be improved by polarization calibration application for certain limitations

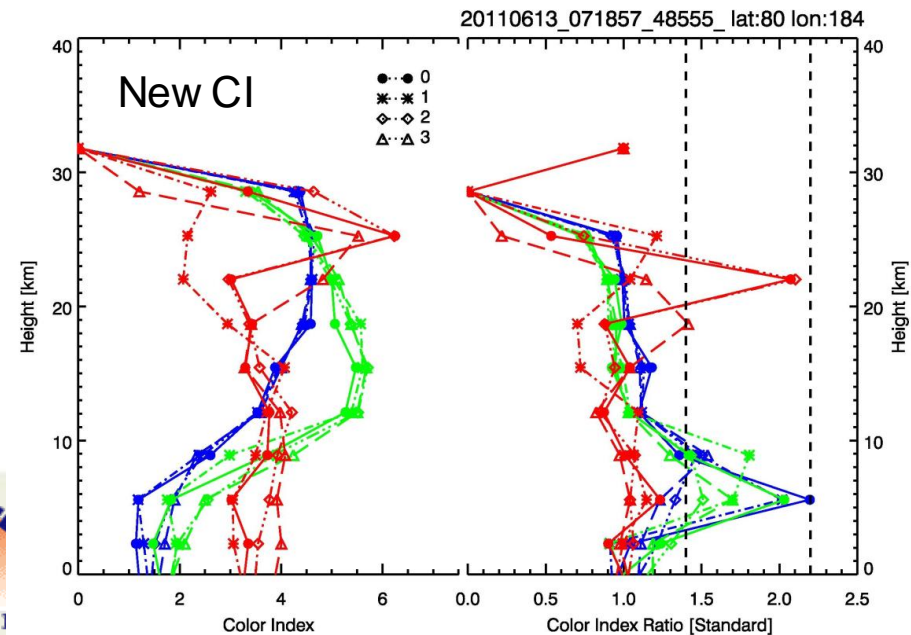
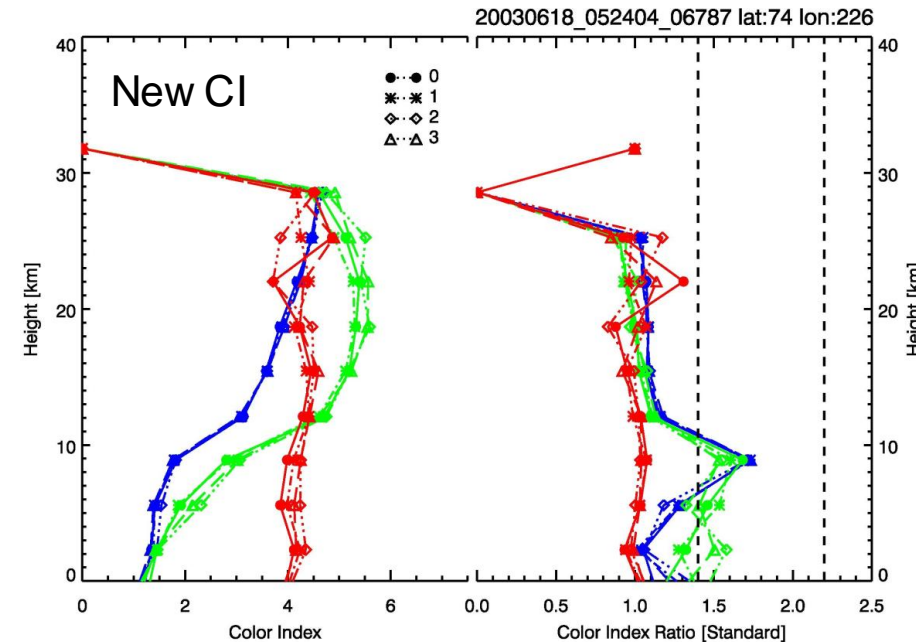
Investigations ongoing:

- Sensivity of Ratios (profiles)
- Threshold optimization
- Ratios with out Rayleigh normalization
- To try the ratio 1050 & 1250 (equidistant from water vapour absorption band)

1090 & 1050

750 & 1090

1550 & 1670



Conclusions:

- 750/1090 nm ratio is not able to distinguish clouds and high aerosol load but 1550/1670 nm ratio can
- Without independent retrieval of OD it may be extremely difficult to distinguish water and ice clouds while still separating clouds/aerosols (*needs more effort and time*)
- Additional pitfall:
TH sampling – it is possible to miss geometrically small and/or optically thin clouds
- Time series analysis show strongly time/lat dependence, may be improved to some extent by switching on polarization calibration (*applicable*)
- Switching to CI instead of CIR in the cloud flagging algorithm is straight (thresholds need to be optimized).